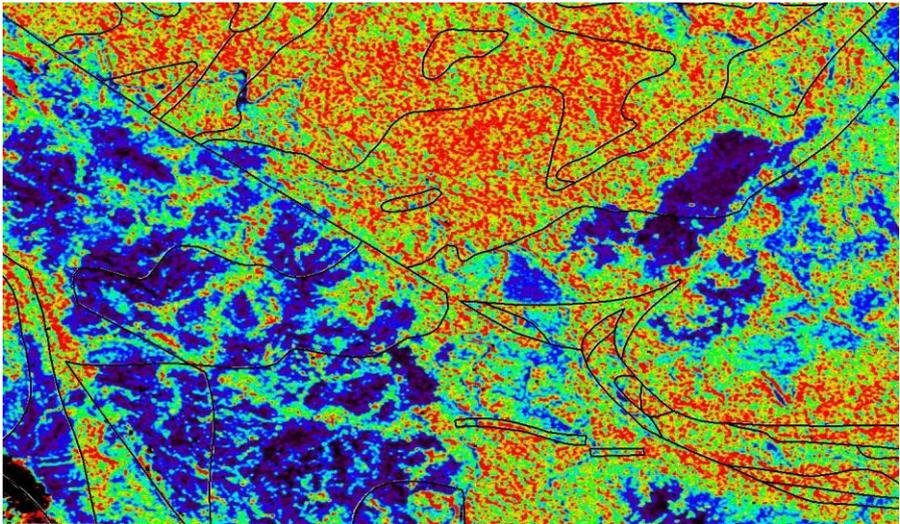
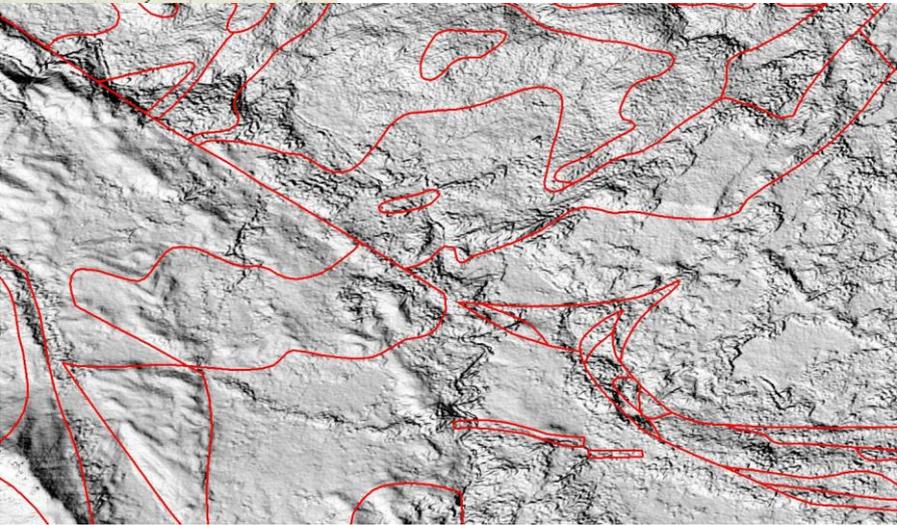
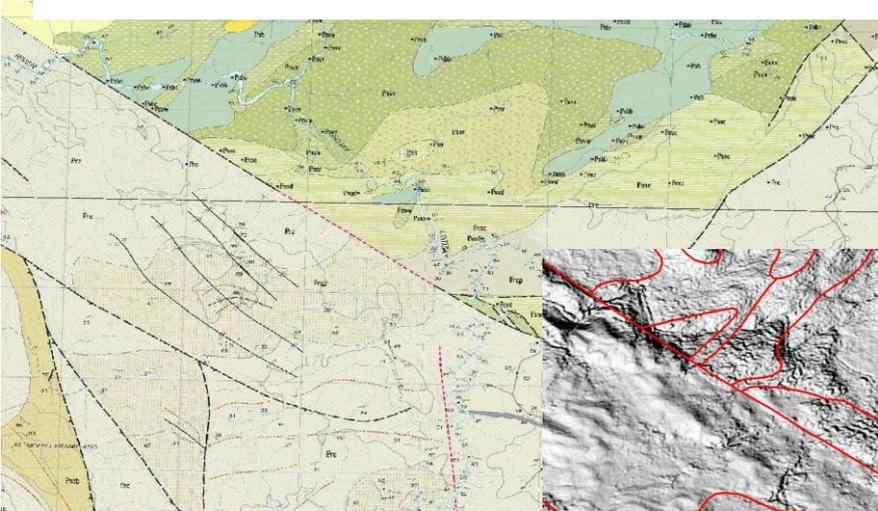


# Application of TOPSAR and Polarimetric SAR for Geological Mapping within north western Tasmania

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# 1 Introduction

## 1.1 Polarimetric radar and TOPSAR basics

Radar remote sensing generates imagery that characterises some of the physical properties (morphology, roughness, dielectric properties and geometric shapes) of the ground surface, including its cover (e.g. vegetation) and near-surface volume (e.g. rock outcrops and regolith layers) (Raney, 1998). It differs from optical or thermal remote sensing in its use of active microwave electromagnetic signals (Figure 1) and interpretation of their backscattered return for electric field strength and polarisation. The operation of radar from either a satellite or airborne platform involves a side looking pulse of microwave radiation transmitted perpendicular to the flight (or orbit) ‘range’ direction and observed by a receiver antenna (Figure 2). A variation of radar systems known as Synthetic Aperture Radar (SAR) optimises the spatial resolution using the Doppler return of the signal along the azimuth flight path of the platform.

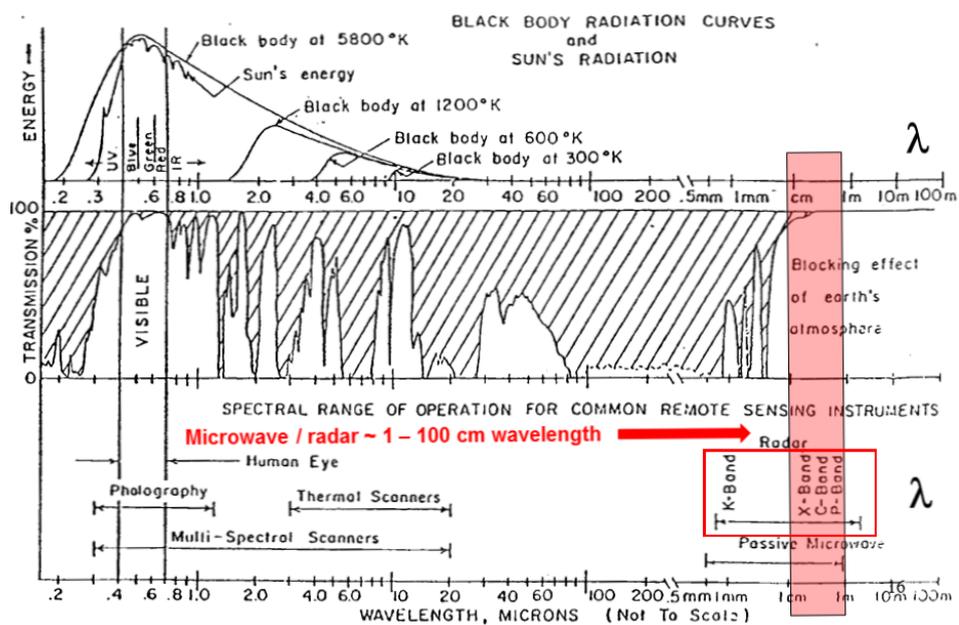


Figure 1 Remote sensing regions including radar in relation to the electromagnetic spectrum. The relative wavelengths of the active radar bands K, X, C and P are shown. Note L band has a wavelength between C and P bands.

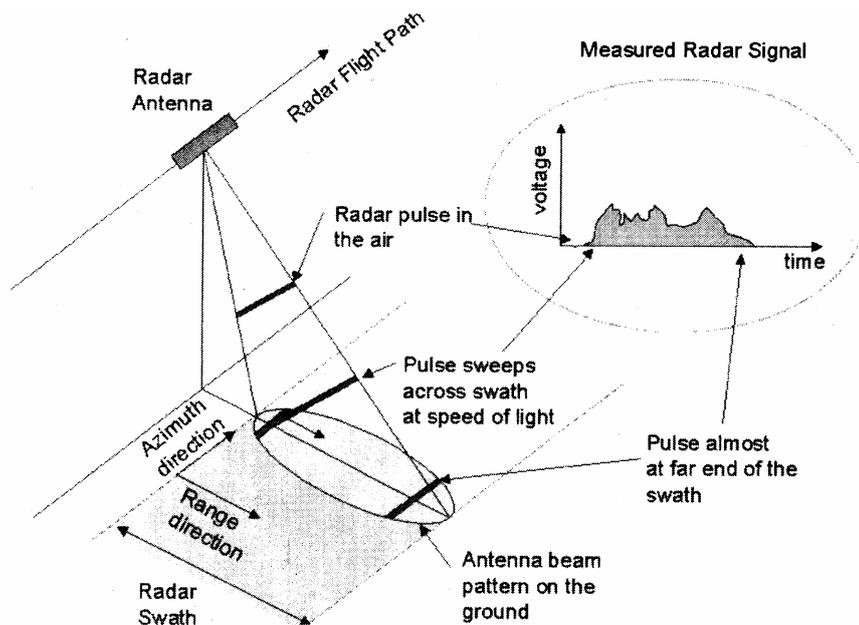


Figure 2 Operation of a radar system

The application radar for the extraction of geological information relies on the surface geometry and its contrasting dielectric properties, typically related to moisture content or mineral composition (Ford, 1998). The geological interpretation of image enhancements are particularly suited to landform analysis from which geomorphological and geological structural interpretation can be made (Tapley, 2002). The radar technique provides its own illumination enabling observations to be independent of cloud cover, light rain, smoke haze and solar illumination, thus allowing all-time observation through all seasons and in all climatic regions. An important capability of radar is the ability to select the illumination geometry, that is, the incidence and azimuth angles, to highlight geological structure and other diagnostic properties of the terrain. Case studies describing SAR radar surveys within densely forested areas of Indonesia and Papua New Guinea have still shown its ability to map both structural features and erosional characteristics of the underlying rocks (Ford, 1998).

Radar technologies such as NASA's multi wavelength polarimetric AIRSAR systems (Lou, 2002; <http://airsar.jpl.nasa.gov/>) include the longer wavelength P (68 cm) and L (25 cm) bands. AIRSAR has the potential capability to measure backscattered signal returns beneath forest canopies related to surface rock outcrops and sub surface regolith (Ford - Henderson and Lewis, 1998; Evans *et al.*, 1988). The AIRSAR system incorporates multi-polarimetric combinations of transmitted horizontal (H) or vertical (V) electric field signals,

and the detection of returning horizontal (H) or vertical (V) electric field returned backscatter. This facilitates the modelling of surface and volume radar scattering associated with vegetation and/or regolith (Figure 3). For example, surface scattering is more likely associated with like-polarized transmitted and returned signals (e.g. HH or VV) while volume scattering is associated with cross-polarized transmitted and returned signals (e.g. HV) (Figure 3). A complete multi-wavelength-polarimetric AIRSAR data set provides all these combinations for C, L and P bands including a total power backscatter measurement (TP). Note that a HV return is assumed to be equivalent to VH polarisation. E.g. C-HH, C-VV, C-HV, C-TP, L-HH, L-VV, L-HV, L-TP, P-HH, P-VV, P-HV, P-TP. Generally the bands associated with longer AIRSAR radar wavelengths, P and to a lesser extent L, measure more backscatter from beneath the vegetation canopy and ground surface and or regolith (Tapley, 2002). In particular, the HV polarization for P (P-HV) band provides the best indication of volume scattering from the shallow subsurface associated with the rock outcrops and regolith. In forested environments, HH polarisation signals potentially suffer less attenuation from the vertically aligned tree trunks and are more likely to provide information about the physical characteristics of the underlying ground-surface (Tapley, 2002). The cross-polarized HV scattering bands are less dependent on incidence angle, varying perpendicular across the flight line, than the co-polarized (HH and VV) backscatter (Tapley, 2002). Consequently HV returns are also less sensitive to variations in terrain slope. Geobotanical relationships, between geological units and vegetation species or forest type are potentially highlighted by the VV polarisation data which is associated with increased interaction with the vegetation and tree trunks (Tapley, 2002).

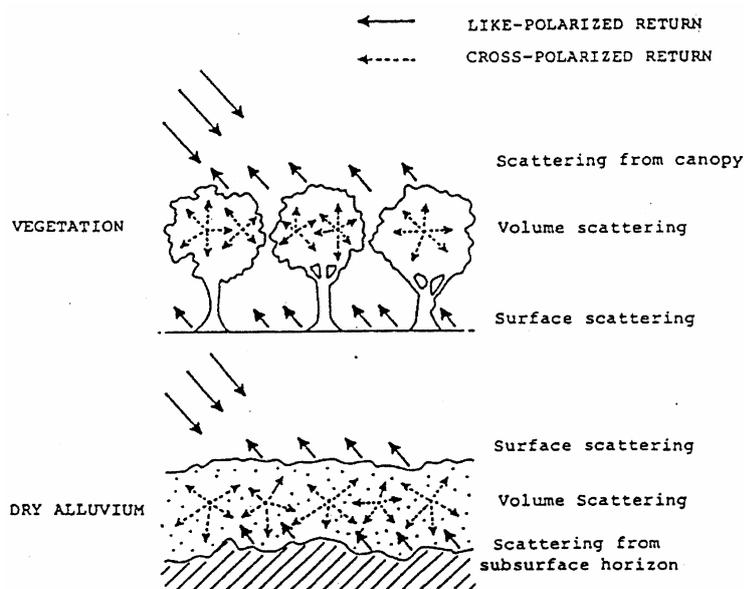


Figure 3 Volume and surface scattering modes for transmitted and returned radar.

In addition, AIRSAR's C band (5.7 cm) is utilised within its TOPSAR system for interferometric topographic mapping (Zebker *et al.*, 1992; Madsen *et al.*, 1993) (Figure 4 a & b). Observing the returned radar backscatter of the C band signal with two detectors, A1 and A2 (Figure 4a) enables the determination of height,  $h$ , for each scanned pixel,  $z(x)$  (Figure 4b). As TOPSAR uses the C-VV wavelength and polarisation, its returned backscatter signals are typically biased by the vegetation canopy.

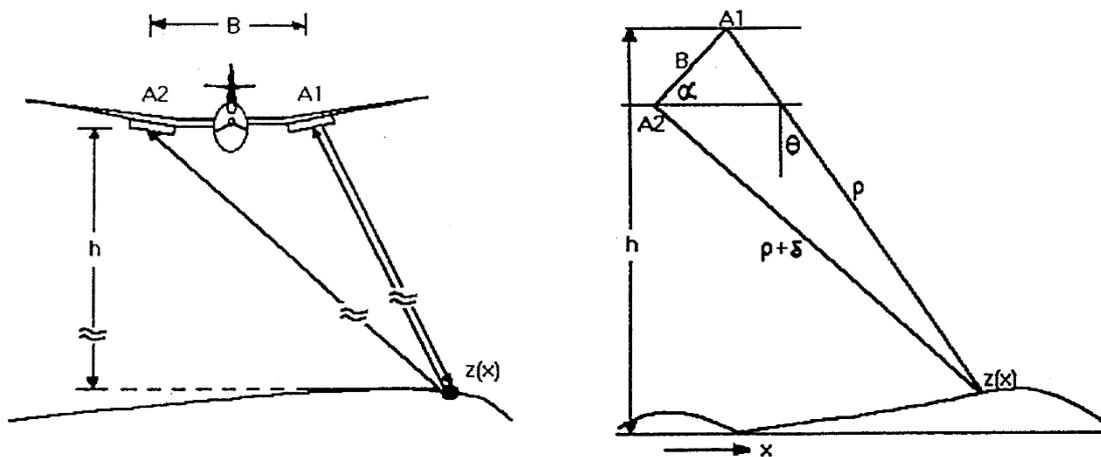


Figure 4 a) Schematic diagram of TOPSAR operation, and b) its corresponding geometry.

## 1.2 Objectives

The aim of this report is to describe the processing and preliminary geological interpretation of TOPSAR acquired by Mineral Resources Tasmania from three flight lines in north western Tasmania during the PACRIM 2 Campaign, August 2000 (Figure 5). As part of this study, the quality control issues of the resulting Digital Elevation Model (DEM) and its comparisons with supplied LiDAR, SRTM 25 metre and 1:25000 Photogrammetric DEM products are examined. The fully polarimetric AIRSAR/POLSAR Mt Read data, acquired in 1993 (Figure 5), was not able to be processed with the current ENVI™ software (<http://www.exelisvis.com/ProductsServices/ENVIProducts/ENVI.aspx>). Its data format appears to be in an atypical format and its processing and interpretation awaits an update by the suppliers of ENVI™.

Mt Read 1993 and Pacrim 2000 AIRSAR scenes

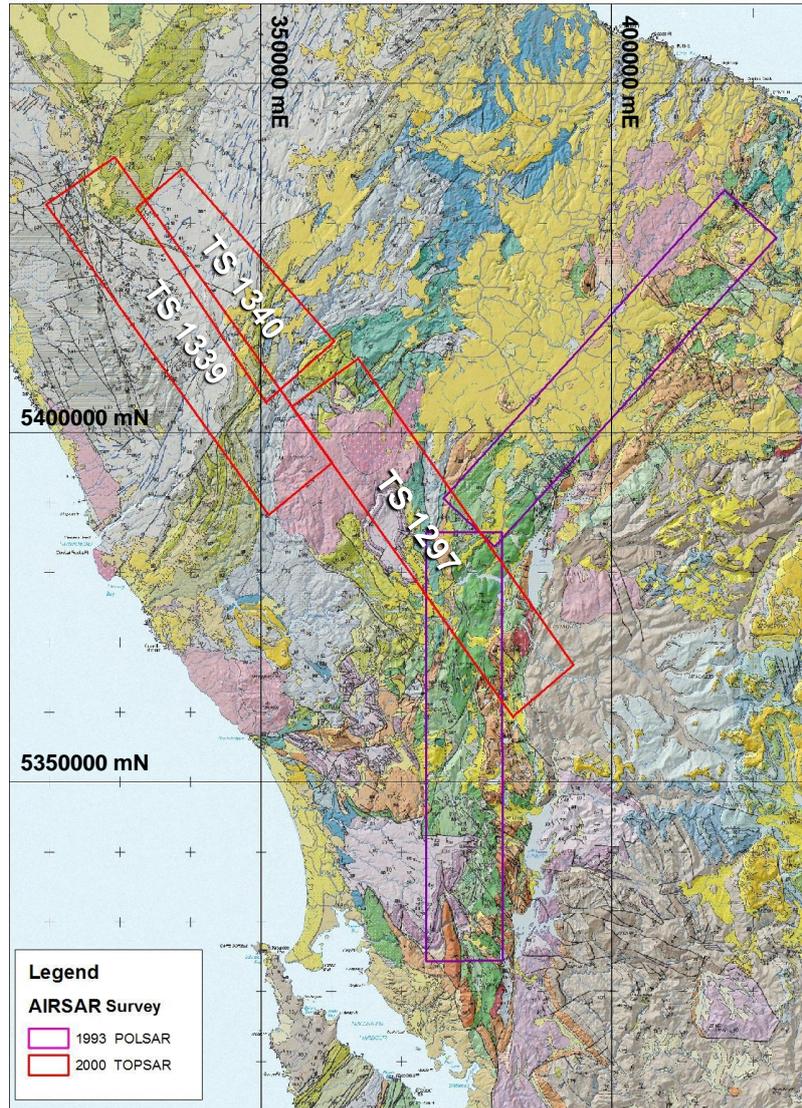


Figure 5 Location of 1993 POLSAR (AIRSAR) and 2000 PACRIM TOPSAR within north western Tasmania.

## 2 Datasets, processing and quality control issues

### 2.1 TOPSAR Polarimetric

The PACRIM 2 TOPSAR data comprised nine channels comprising C-VV, L-HH, L-VV, L-HV, L-TP, P-HH, P-VV, P-HV, and P-TP. Three flight lines of TOPSAR imagery, TS1297, TS1340 and TS1339 were acquired of approximately 60 x 10 km, 60 x 10 km and 30 x 10 km image areas respectively (Figure 5). These were imported into ENVI™ software which facilitates the synthesis of the multiple wavelengths and polarisations from the Stokes Matrix Compression format (<http://airsar.jpl.nasa.gov/documents/dataformat.htm>). The nine channel imagery was then registered in a three stage process :

- 1) polarimetric TOPSAR imagery were registered using the WGS 84 geographic coordinates within the supplied header information for the corners (Appendix 1);
- 2) the WGS 84 projected imagery was converted to GDA94 MGA Zone 55S Easting and Northing coordinate projection;
- 3) The GDA94 MGA Zone 55 was registered using 20 to 30 ground control points identified by features (watercourse/road intersections, dam walls, mountain peaks etc) identified on supplied MRT 1:25000 topographic maps, SPOT imagery and 1:25,000, and 1:25000 Photogrammetric and SRTM 25 metre DEMs (Figure A1, Appendix 1).

The resulting GDA94 MGA Zone 55 polarimetric TOPSAR imagery was imported into ERDAS ERMapper™ (<http://erdas-er-mapper.software.informer.com/>) for integration and comparison with supplied MRT datasets. Some issues of systematic spatial noise were apparent and likely to have resulted from AIRSAR system and aircraft electronics (Tapley, 2002). Two dimensional (2D) Fast Fourier Transform (FFT) 2D band cut filtering was attempted using ENVI™ however difficulty was encountered in the exclusion of the systematic spatial frequency noise while not introducing noise from the user defined band cuts. As a consequence, no FFT filtering was applied to this TOPSAR imagery. However a more serious noise issue was apparent from radar “speckle”, associated with inherent random and multiplicative returning coherent backscatter (Tapley, 2002). Various adaptive filtering techniques using both ENVI™ and ERDAS ERMapper™ software including Median, Gamma, Froist and Lee filters (Lopes *et al.*, 1990). The standard Lee filter (Lee, 1980) available using ERDAS ERMapper™ was found to give the best result in reducing the speckle for this AIRSAR/TOPSAR imagery. This operation of this filter is based on the probability of a Gaussian distribution where 95.5% of random samples are assumed to be within a 2 standard deviation (2 sigma) range. This noise suppression filter replaces the pixel of interest with the average of all DN values within a 5 x 5 pixel sample moving window that fall within the designated range (ERDAS, 1999)

## 2.2 TOPSAR DEM

The PACRIM 2 TOPSAR DEM data comprised four bands including the C-VV polarised band, the correlation information (DEM accuracy estimate), the incidence angle (degrees) of the radar signal and the calculated topographic DEM. The conversion from the supplied TOPSAR data format into these four parameters was done using ENVI™. Missing radar returns from radar shadows produced by the topography were replaced using standard surface fitting techniques available with ENVI™ although the interpolation was still apparent

in some areas of these TOPSAR DEM imagery. The resulting four parameters associated with the C band derived DEM comprised identical number of pixel and line samples as the TOPSAR polarimetric file for each flight line and WGS 84 geographic coordinates within the supplied header information (Appendix 1). Consequently the same registration procedure listed in Section 2.1 was followed to generate equivalent GDA94 MGA Zone 55S imagery for the three flight lines.

## 2.3 Issues affecting radar interpretation and quality control

### 2.3.1 Incident angle

The sensitivity of a radar systems response to the geometry of its operation and the acquisition of returning backscatter is illustrated by Figures 6 and 7. In particular Figure 6 shows the issue of layover distorting the dimensions of the topographic slopes for large radar depression angles, while radar shadowing can be present for shallow depression angles. Typically the AIRSAR/TOPSAR system describe their acquisition specifications in terms of look angle, which is related to the depression (*e.g.* =  $90 - \text{look angle degrees}$ ) and incident angles as shown in Figure 7a). In a flat ground surface situation, the look angle and incident angles are equal (Figure 7a) however the local incident angle can be significantly affected by local topography (Figure 7b). When the local incidence angle is smaller than the local slope angle layover effects and distortion will be apparent (Raney, 1998).

The high spatial sampling resolution of 5 metres acquired by the AIRSAR/TOPSAR imagery makes the local incident angle particularly important in the steeply dipping relief of western Tasmania. The range of look angles measured by the AIRSAR/TOPSAR PACRIM2 surveys varied between approximately 25 and 64 degrees. In general the larger look angles acquired by AIRSAR/TOPSAR at the furthest extent of the swath or range (*e.g.* 50-64 degrees) will generally accentuate geomorphic and structural features related to topographic relief and outcrop by their apparent shadowing effect. At shallower look angles (*e.g.* 25-35 degrees) a greater chance of spatial distortion and layover effects is also possible, depending on the local relief.

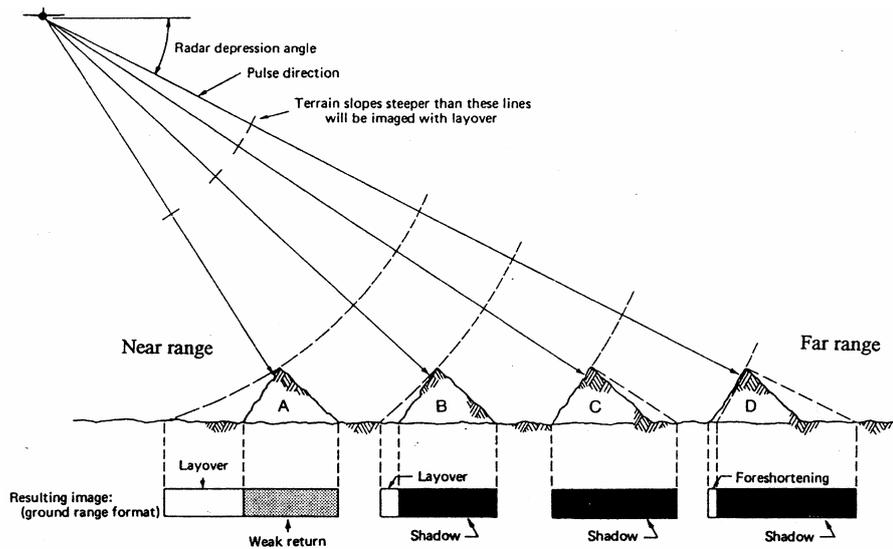
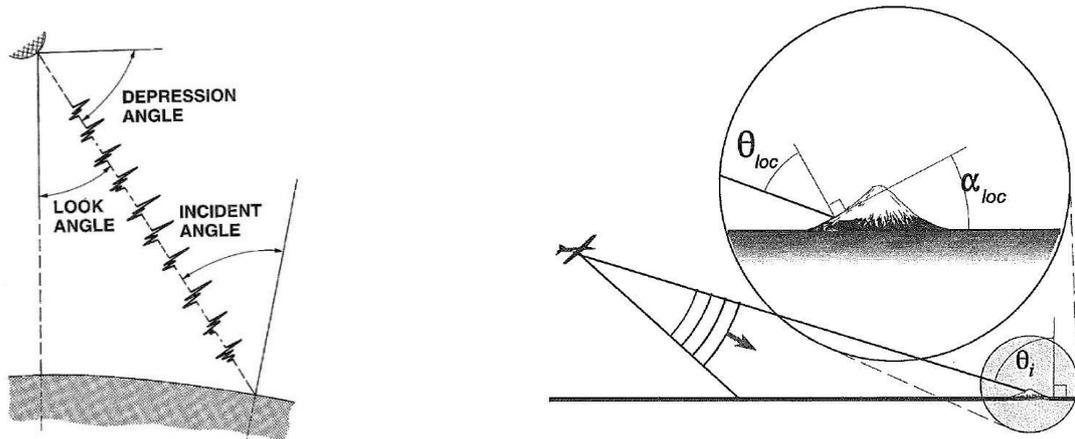


Figure 6 Schematic diagram illustrating the interaction between topography and the radar depression angle on the resulting radar image effects of layover and shadowing (Elachi & Van Zyl, 2006).



a)

b)

Figure 7 a) Relationship between look angle, depression angle and incident angle assuming a relatively flat ground surface (Ford, 1998); b) Schematic diagram illustrating local incident angle  $\theta_{loc}$  and the average incident angle  $\theta_i$  resulting from local slope angle  $\alpha_{loc}$  (Raney, 1998).

### 2.3.2 Surface roughness

In general radar backscatter increases with increasing surface roughness depending on the radar wavelength (Raney, 1998). Geological applications of AIRSAR/TOPSAR utilise these effects of the differing responses to surface roughness by observing the response to the three different band wavelengths. In particular, the surface roughness properties of the rock outcrops and regolith, will generate variable amounts of diffuse backscatter according to the wavelength of the radar pulse and its incident (local) angle. Surface roughness is typically measured in terms of root mean square (rms) measures varying from millimetres , centimetres to metres for very rocky terrain. Quantitatively, a surface is considered definitely “rough” (e.g. generating a high backscatter response) if :

$$h > \lambda / (4.4 \text{ Cos } \theta) \quad (1)$$

where  $h$  : rms height,  $\lambda$  : radar wavelength, and  $\theta$  : incident angle (Ford, 1998).

Alternatively a surface is considered definitely “smooth” (e.g. generating a low backscatter response) if :

$$h < \lambda / (25 \text{ Cos } \theta) \quad (2)$$

(Ford, 1998).

Assuming the AIRSAR/TOPSAR system parameters for these data sets, the following rms height definitions for “rough” and “smooth” surface roughness were calculated using Eq (1) and Eq. (2) for the minimum and maximum look angles across the radar acquisition swath. No allowance is made for topographic relief and the incident angle,  $\theta_i$ , is assumed to be equivalent to  $\theta_{loc}$  and the survey look angle (Figures 7a&b).

Table 1 highlights the potential variation of the AIRSAR/TOPSAR response to look angle across the swath of the imagery (Figure 2) without taking the topographic effects into account. P, and to a lesser extent L radar band, are more likely to observe a “smooth” radar surface at the far swath extent or range of the imagery while C band is most likely to register rough radar surfaces at the near swath range (Table 1). The P band radar imagery is also most likely to discriminate the roughest surface nature at the nearest swath extent of the imagery. The question of whether the surface is a thick vegetation canopy or rocky ground surface is also affected on the ability of the longer wavelength radar to penetrate vegetation and also its signature polarisation. As discussed in Section 1.1 and illustrated in Figure 3, volume scattering measured by P band (and to a lesser extent L band) with a HV polarisation provides the optimal imagery to observe geologically related surface roughness

beneath vegetation canopy. HV is also optimal for minimising the effects of variable local incident and look angles (Tapley, 2002).

<b>AIRSAR/TOPSAR Band</b>	<b>C</b>	<b>L</b>	<b>P</b>
<b>Wavelength (<math>\lambda</math>, cm)</b>	<b>5.7</b>	<b>25</b>	<b>68</b>
<b>Response at far swath range :</b>			
"Rough" rms height (cm) @ $\theta_i = 68$ degrees	>3.5	>15.2	>41.3
"Smooth" rms height (cm) @ $\theta_i = 68$ degrees	<0.6	<2.7	<7.3
<b>Response at near swath range :</b>			
"Rough" rms height (cm) @ $\theta_i = 25$ degrees	>1.4	>6.3	>17.1
"Smooth" rms height (cm) @ $\theta_i = 25$ degrees	<0.3	<1.1	<3.0

Table 1 Definitions of smooth and rough radar surfaces for AIRSAR/TOPSAR radar band wavelengths and maximum-minimum range of look angles where Look angle = Incident angle, assuming a flat ground surface.

Studies by Huadong *et al.*, (1996) using shuttleborne C and L band multi-polarisation radar found that L band HV polarised imagery was the best combination for discriminating two different basalt units, alluvium and bedrock. Kierein-Young *et al.* (1992) showed that quantifiable estimates of surface roughness rms is possible in an arid environment such as Death Valley, Nevada. Height rms values were calculated from modelling the fractal dimensions of the C, L and P band power spectral density however the model assumed generally smooth playa to partly smooth alluvial surfaces (Kierein-Young *et al.*, 1992). Subsequent communication with one of the authors, indicated that in thickly forested environments there are wavelength-specific problems related to partial canopy penetration, leaf moisture dielectric effects and canopy/understory vegetation cover that combine to make this a difficult problem (Kruse, pers. comm.). The observation of surface roughness is also identified by Raney (1998) to be more subtle and less quantifiable with increasing penetration through vegetation cover. Ford (1998) also notes that estimation of surface roughness using multi wavelength radar can be possible within sparsely vegetated terrains.

### 2.2.3 Dielectric properties and Moisture

Another factor that can effects radar backscatter is the dielectric property and related moisture content of the surface materials. Such dielectric properties influence the attenuation and reflectance of the electric field component of the radar pulse. Most rocks exhibit a narrow range in dielectric properties however they will vary significantly as a function of water moisture, its porosity and water holding capacity (Ford, 1998). Water can strongly limit

the ground penetration of even the longer wavelength L and P band radar when interacting with moist soil and regolith (Ford, 1998) (Figure 8). As a consequence the radar backscatter response can vary between geological units when they contain different weathered regolith profiles affecting their porosity and moisture content.

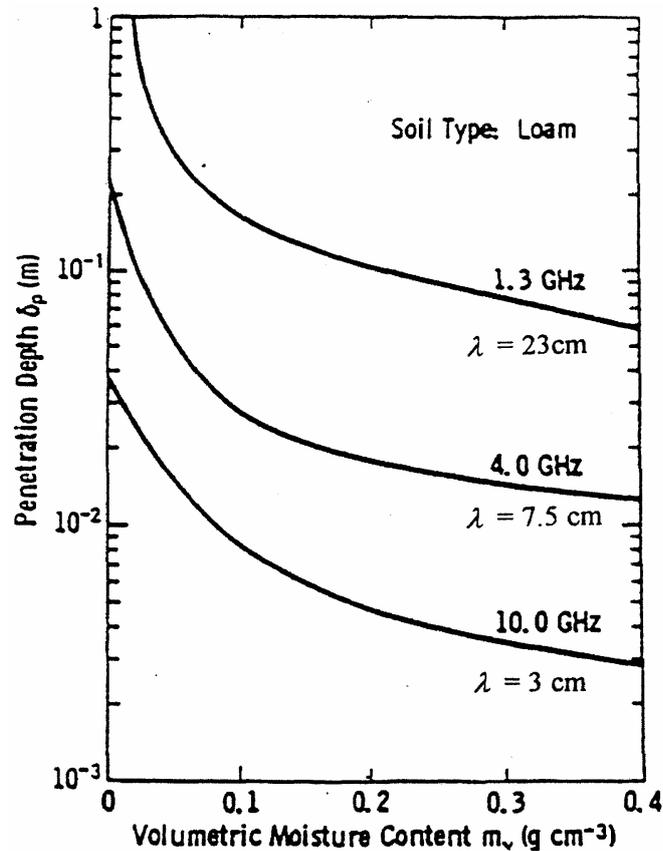


Figure 8 Penetration depths for soil with different moisture contents at different wavelengths e.g. C band ( $\lambda \sim 3\text{cm}$ ) and L band ( $\lambda \sim 23\text{cm}$ ), after Dobson (1993).

### 2.3.4 Combined QC issues

The effect of the variable incident angles across each individual flight swaths upon the returned backscatter limits the ability to simply mosaic radar imagery, particularly if flights are in opposite look angle directions (e.g. TS1339 acquired with NE range versus the SW range acquired TS1340 and TS1297). Consequently the presence of radar shadowing and layover will change according to the topographic relief (Section 2.3.1). Such variable look and hence incident angles will also significantly affect the response of surface roughness across the swath for each AIRSAR wavelength (Section 2.3.2). Changes in vegetation and/or soil moisture due to precipitation between flights could also affect radar backscatter however the acquisition of all three flights on 17/8/2000 makes this a more unlikely influence

of image boundary differences. Due to the complexity of these various effects, no attempt has been made to level the radar products.

## 3 Results

### 3.1 TOPSAR DEM

#### 3.1.1 Image products

The TOPSAR DEM product processed from the three Tasmanian flight lines appeared to generate good quality data at 5 metre spatial sampling that mosaicked well as a combined DEM data set encompassing 1440 km<sup>2</sup> and a range of elevation between 3 to 1370 metres. Examples of the mosaicked DEM product as an image, as a shaded relief (illuminated at 45 degree NE and 60 elevation) and a topographic modelled slope, are shown in Figures 9a, 9b and 9c respectively.

Associated with the supplied TOPSAR DEM product are imagery representing estimates of the local incidence angle (0-180 degrees) to assist in the interpretation of the radar acquisition. Each acquisition will generate a specific incidence angle product relevant for the flight line track and look direction. Figure 10 shows the example of the incidence angle estimates for flight line, TS1340. The lower incident angles shown along the north eastern areas of the imagery generally correspond to the lower radar look angles operating by the AIRSAR/TOPSAR as it flew along a flight path just beyond the NE of the imagery (Figure 10). Although the range of look angle available by the survey is restricted by the system specification, the effective incidence angle will be dependent according to the local incident angle as shown in Figure 7b. As discussed previously, such areas dominated by lower incident angles will be more susceptible to layover distortion while radar shadow effects will increase at the higher incident angles (Figure 10).

A statistical comparison between the overlapping DEM results for the TOPSAR acquisitions (Figure 11) was undertaken to evaluate the consistency of the derived elevations and registration procedure (Appendix 2). Although there appeared spurious variations for the minimum and maximum elevations values, the correlation coefficients were greater than 0.99, and the median / mean values showed a less than 5 metres difference (Appendix 2).

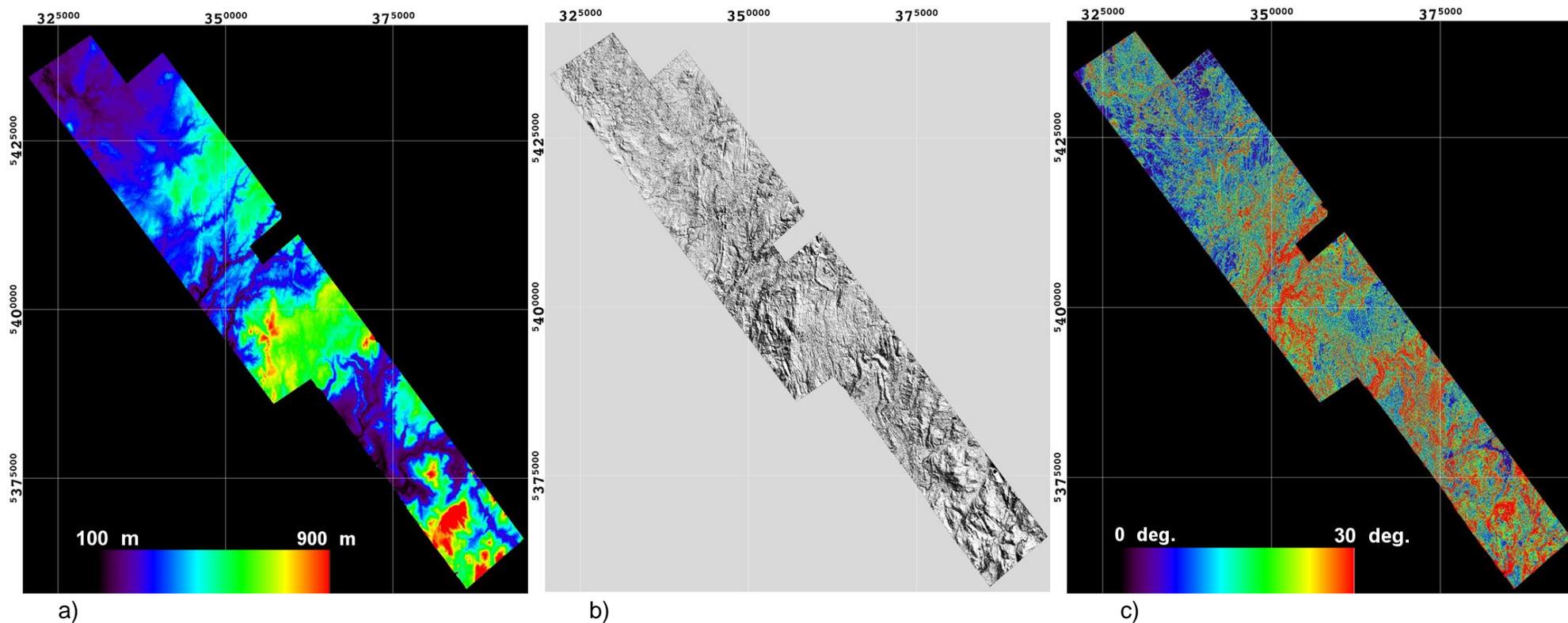


Figure 9 a) TOPSAR DEM mosaic of TS1297, TS1339 and TS1340 acquisitions displayed with a linear 100 to 900 metre elevation stretch; b) Shaded mosaic illuminated with 45° NE azimuth and 60° elevation; and c) Slope mosaic product displayed with a linear 0 to 30 degrees slope stretch. All imagery is registered for GDA94 MGA Zone 55

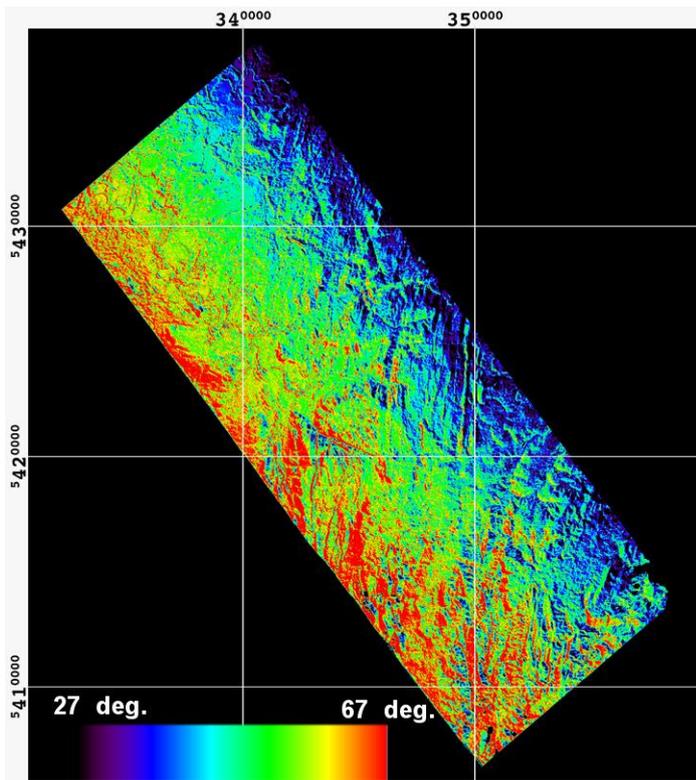


Figure 10 Supplied calculated local incident angle estimates for flight line TS1340.

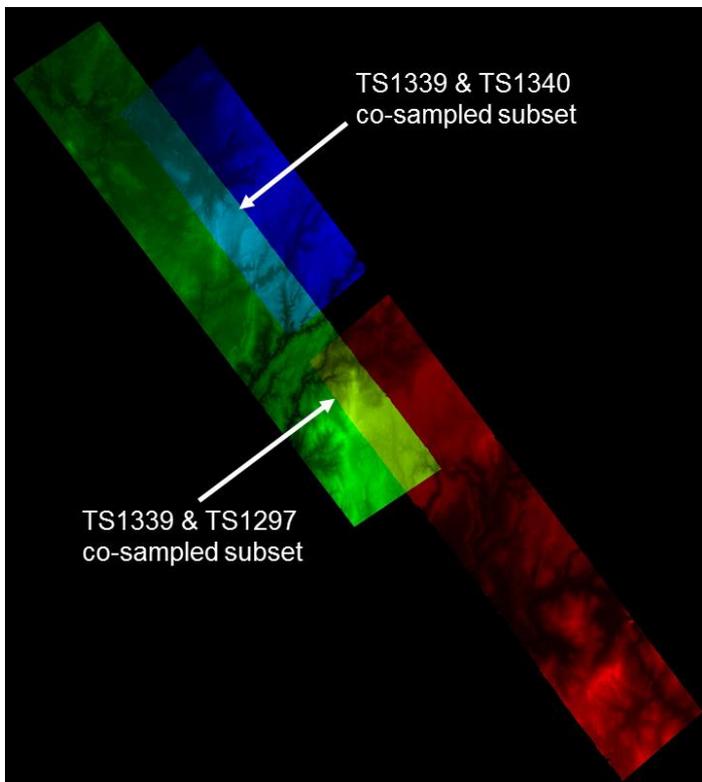


Figure 11 Depiction of the three TOPSAR acquisitions and their overlapping areas. The spatial extent is the same as shown in Figure 9 a-c).

### 3.1.2 Comparison with SRTM 25m, Photogrammetric derived and LiDar DEMs

Statistical comparisons between the mosaicked TOPSAR DEM product and the MRT supplied SRTM 25 metre, photogrammetric derived and LiDar DEMs was undertaken to assess its quality. Issues of variable spatial sampling resolution and registration between these four datasets is potentially relevant however it was still seen as a useful exercise for validation. ERDAS ERMapper™ software was used to calculate the statistics and subset according to their spatial extents. Table 2 shows the comparison between the TOPSAR SRTM 25 metre, photogrammetric derived DEMs, while Table 3 shows the comparison for the significantly smaller coincident area encompassing all four DEMs including SRTM 25 metre, photogrammetric and LiDar derived DEMs. The relevant LiDar data sets for this comparison were the Meredith, South1 and North1 acquisitions. The statistics for the two coincident areas showed that although the TOPSAR DEM exhibited significantly lower minimum values than the other DEMs, the overall comparison was favourable with a correlation of 0.996 and higher (Table 2 and 3).

	<b>TOPSAR</b>	<b>SRTM 25m</b>	<b>Photogrammetry</b>
<b>Area (Hectares)</b>	144029	144029	144029
<b>Spatial sampling (m)</b>	5	25	25
<b>Minimum</b>	3.1	54.6	50.0
<b>Maximum</b>	1369.8	1246.8	1266.6
<b>Mean</b>	365.1	364.2	361.4
<b>Median</b>	334.1	329.4	325.6
<b>Std.Dev.</b>	163.4	164.0	165.9
<b>CorrelationMatrix</b>	<b>TOPSAR</b>	<b>SRTM 25m</b>	<b>Photogrammetry</b>
<b>TOPSAR</b>	1	0.997	0.996
<b>SRTM25m</b>	0.997	1	0.999
<b>Photogrammetry</b>	0.996	0.999	1

Table 2 Statistical comparison between coincident TOPSAR, SRTM 25m and Photogrammetry derived DEMs.

	<b>TOPSAR</b>	<b>SRTM 25m</b>	<b>Photogrammetry</b>	<b>LiDAR</b>
<b>Area (Hectares)</b>	17653	17653	17653	17653
<b>Spatial sampling (m)</b>	5	25	25	0.5 - 2.0
<b>Minimum</b>	3.5	102.6	97.0	95.0
<b>Maximum</b>	917.0	915.8	920.0	1160.6
<b>Mean</b>	434.5	437.1	434.7	434.0
<b>Median</b>	481.7	483.8	479.6	478.0
<b>Std.Dev.</b>	166.3	162.9	164.8	164.3
<b>CorrelationMatrix</b>	<b>TOPSAR</b>	<b>SRTM 25m</b>	<b>Photogrammetry</b>	<b>LiDAR</b>
<b>TOPSAR</b>	1	0.998	0.998	0.996
<b>SRTM25m</b>	0.998	1	0.999	0.998
<b>Photogrammetry</b>	0.998	0.999	1	0.998
<b>LiDAR</b>	0.996	0.998	0.998	1

Table 3 Statistical comparison between coincident TOPSAR, SRTM 25m, Photogrammetry and LiDar derived DEMs.

A further analysis was undertaken to assess the accuracy of the elevations and spatial Easting and Northing coordinates of the TOPSAR, relative to the other supplied DEMs and, in particular, to the 1:25,000 topographic mapping. The effect of the geology on the derived DEM elevation values was also briefly examined. In particular, two statistical training subsets were chosen from the areas with coincident DEM coverage, one encompassing the Meredith intrusives (Dg, Dga), and another with no Meredith intrusives, dominated by the Mt Read volcanics (Figure 12). The statistical results for each of the two training areas and the entire coincident DEM area are listed in Appendix 3. There again appeared differences between the minimum and maximum DEM values from spurious outliers however the mean and median values were fairly similar within both the Meredith and non Meredith granitic areas. The Meredith Granite area showed a smaller variation in median elevations than the volcanic terrain (e.g. 3.82 m cf 7.06 m) (Appendix 3). The correlation coefficients were generally high for both areas although the volcanic terrain showed higher values (> 0.998) compared to the Meredith granite (> 0.95). However a spurious LiDar maximum of 1160 metres possibly generated this difference.

A detailed comparison between the DEM data sets was also undertaken using five spot topographic peaks within both training areas (Table 4). The supplied 1:25,000 topographic mapping with surveyed elevations were used as a reference to calculate the difference (“ $\Delta(1:25,000)$ ”) relative to the four DEM data sets (Table 4). No significant differences were observed and generally within 50 metres coordinate accuracy and better than 15 metres elevation accuracy. There was potentially subjectivity in choosing the position of the

topographic peaks from the various DEMs which may explain the discrepancy of the Mt Sale Lidar result (Table 4). These results appeared consistent with estimates of spatial x-y registration errors suggested by the Root Mean Square (rms) values obtained for the first order polynomial transformation for each individual TOPSAR flight line (Appendix 1). In particular, rms values ranged from between 4.5 and 9.4 pixels (e.g. 24 to 47 metres) (Appendix 1).

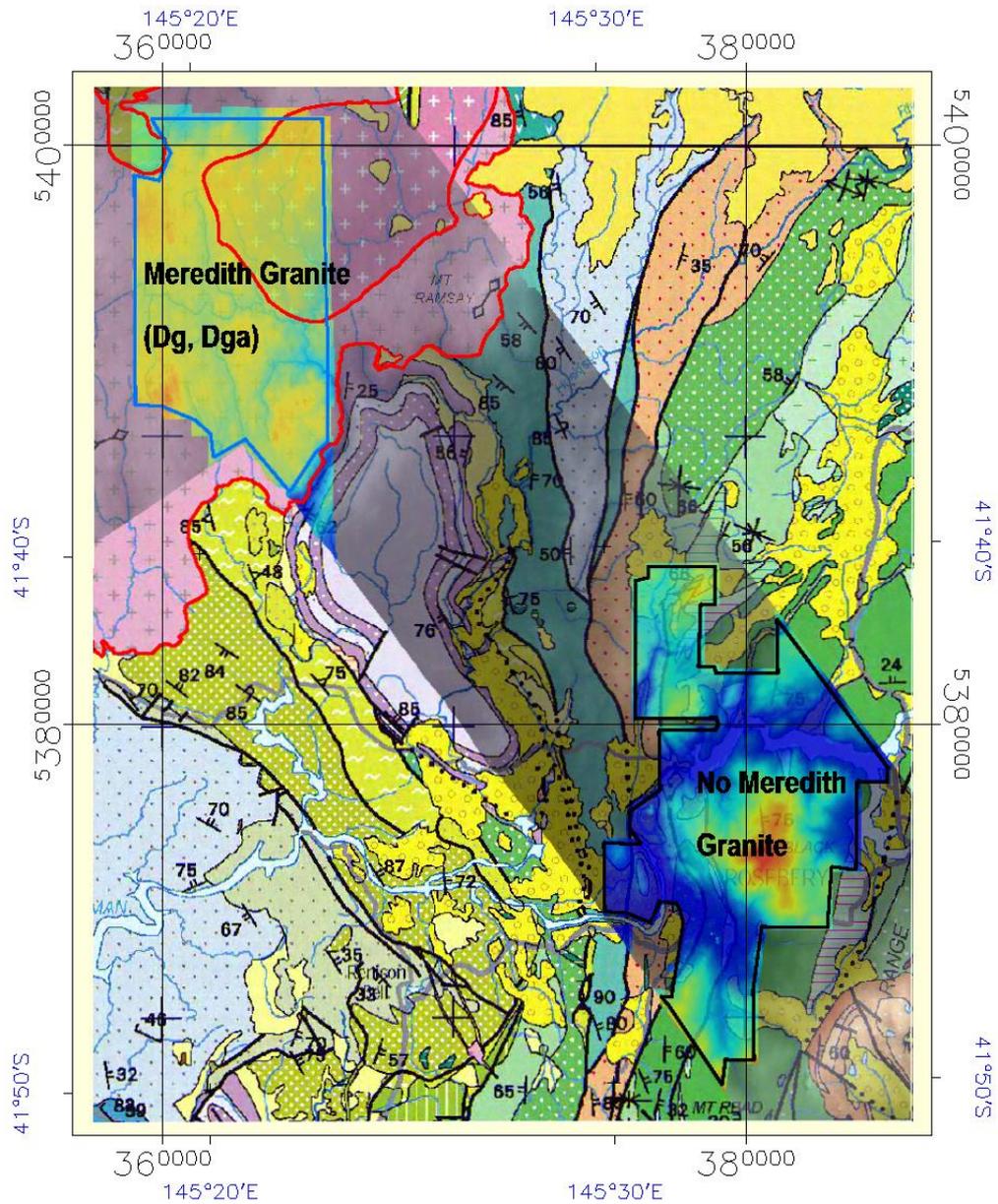


Figure 12 Areas of coincident TOPSAR (blue:low; red:high), SRTM 25m, Photogrammetric 25m and LiDAR DEMs overlaying 1:250,000 geology. The Meredith Granite (DG, Dga – red boundary, blue subset) and non-Meredith granite (black subset) highlight the two training areas for statistical comparison. The mosaicked TOPSAR DEM survey area is shown in grey.

Coordinates (GDA94)	Mt Black	$\Delta(1:25,000)$	Mt Sale	$\Delta(1:25,000)$	Burns Peak	$\Delta(1:25,000)$		$\Delta(1:25,000)$		$\Delta(1:25,000)$
1:250,000 Geology:	Edv		Edv		Edsvl		Dg/Dga		Dg/Dga	
(Easting) 1:25,000	380647		382944		378385		363972		362842	
(Northing) 1:25,000	5375580		5377974		5384652		5394495		5398980	
(Elev.) 1:25,000	929		521		660		584		623	
(Easting) TOPSAR	380642	-5	382932	-12	378373	-12	363963	-9	362839	-3
(Northing) TOPSAR	5375598	18	5377978	4	5384656	4	5394546	51	5399011	31
(Elev.) TOPSAR	916	-13	532	11	652	-8	588	4	621	-2
(Easting) SRTM25m	380667	20	382895	-49	378395	10	363995	23	362867	25
(Northing) SRTM25m	5375604	24	5377940	-34	5384678	26	5394451	-44	5398991	11
(Elev.) SRTM25m	916	-13	525	4	653	-7	572	-12	612	-11
(Easting) Photogram.	380626	-21	382923	-21	378363	-22	363977	5	362839	-3
(Northing) Photogram.	5375611	31	5377980	6	5384652	0	5394498	3	5398991	11
(Elev.) Photogram.	920	-9	520	-1	650	-10	580	-4	620	-3
(Easting) LiDAR	380644	-3	382848	-96	378401	16	363991	19	362857	15
(Northing) LiDAR	5375591	11	5377960	-14	5384637	-15	5394510	15	5398970	-10
(Elev.) LiDAR	927	-2	528	7	662	2	582	-2	621	-2

Edv : Felsic to intermediate calc-alkaline volcanic rocks, (Central Volcanic Complex and correlates / Mt Read Volcanics)

Edsvl : Felsic lava within Western Volcano-Sedimentary Sequence and correlates

Dg/Dga : Meredith granite and associated intrusives

Table 4 Spot comparisons between the TOPSAR, SRTM25m, Photogrammetric 25m and Lidar DEMs with five 1:25,000 topographic located hilltops.

### 3.2 TOPSAR multi-wavelength and multi-polarimetric products

The image products from the three AIRSAR/TOPSAR acquisitions were imported into an ERDAS ERMapper™ algorithm style GIS database incorporating ancillary MRT supplied vector and raster files of other DEMs, SPOT imagery, orthophotography, geology, vegetation, and airborne geophysics. Two areas within TS1340 were studied in particular in order to observe the effects of incident angle, wavelength and polarisation as well as compare with published geological and vegetation information (Figure 13). In particular it was attempted to examine the issues of the potentially competing backscattering effects of vegetation versus geological. Figure 13 also shows the enhanced backscatter of the C band VV polarisation (Cvv) near swath range along the north eastern edge of the image associated with the lower incident angle.

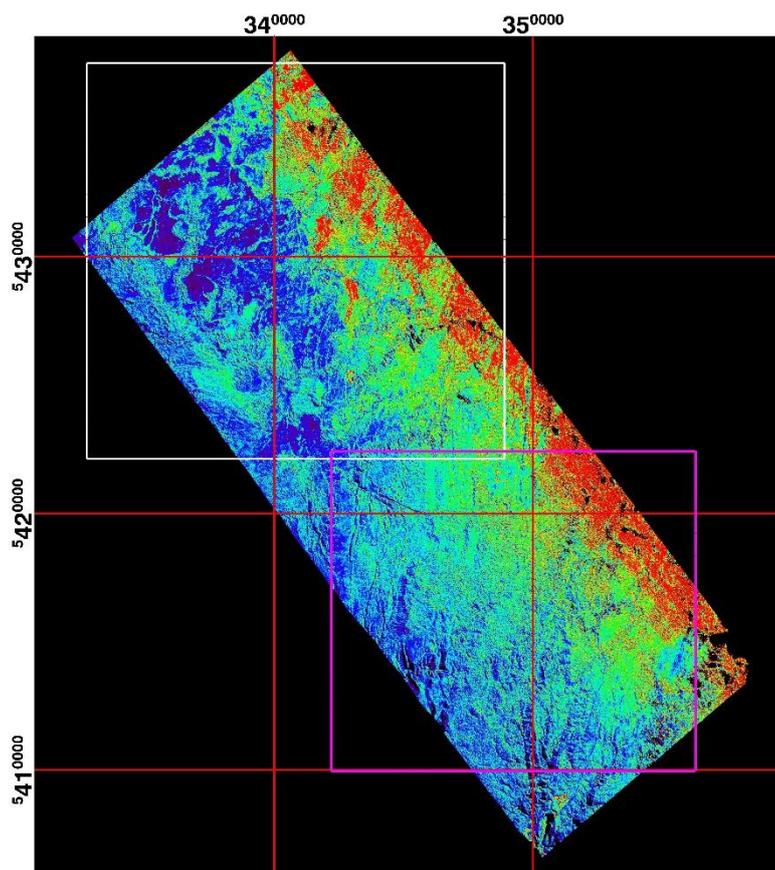
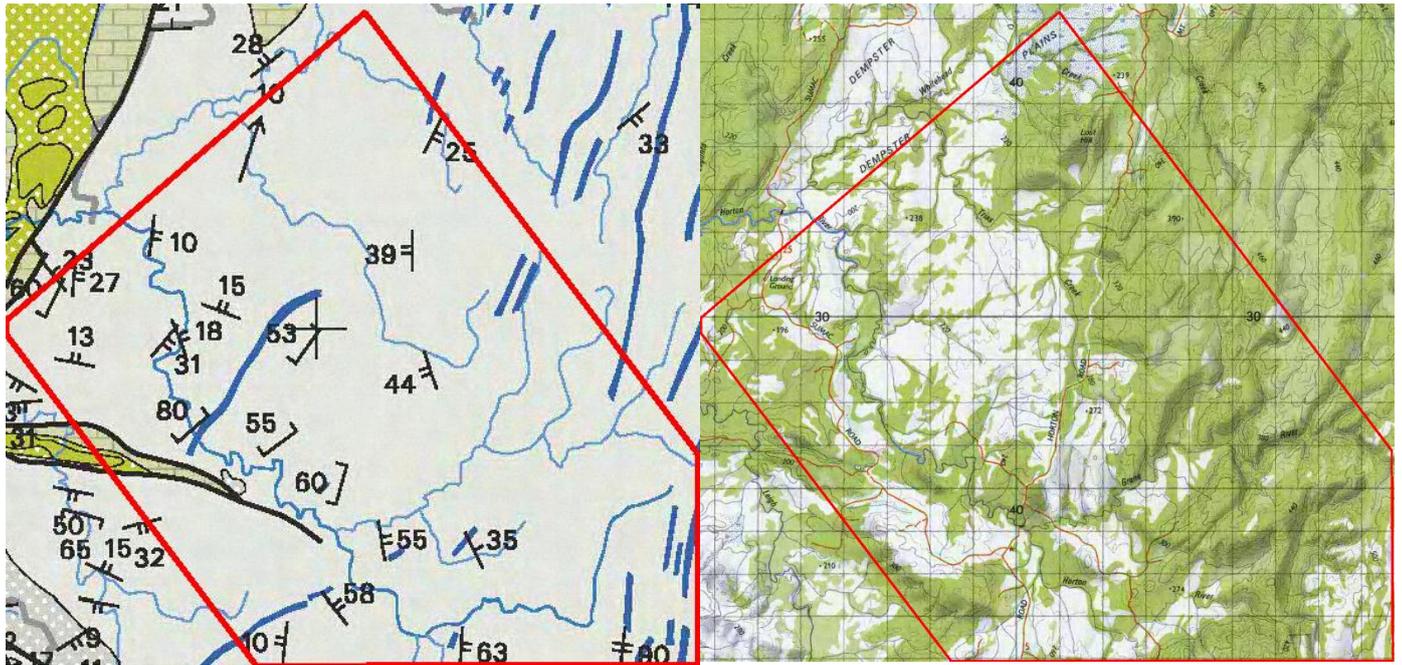


Figure 13 AIRSAR/TOPSAR Cvv image of flight line TS1340 showing the locations of the vegetation (white box) and geological (magenta box) study areas.

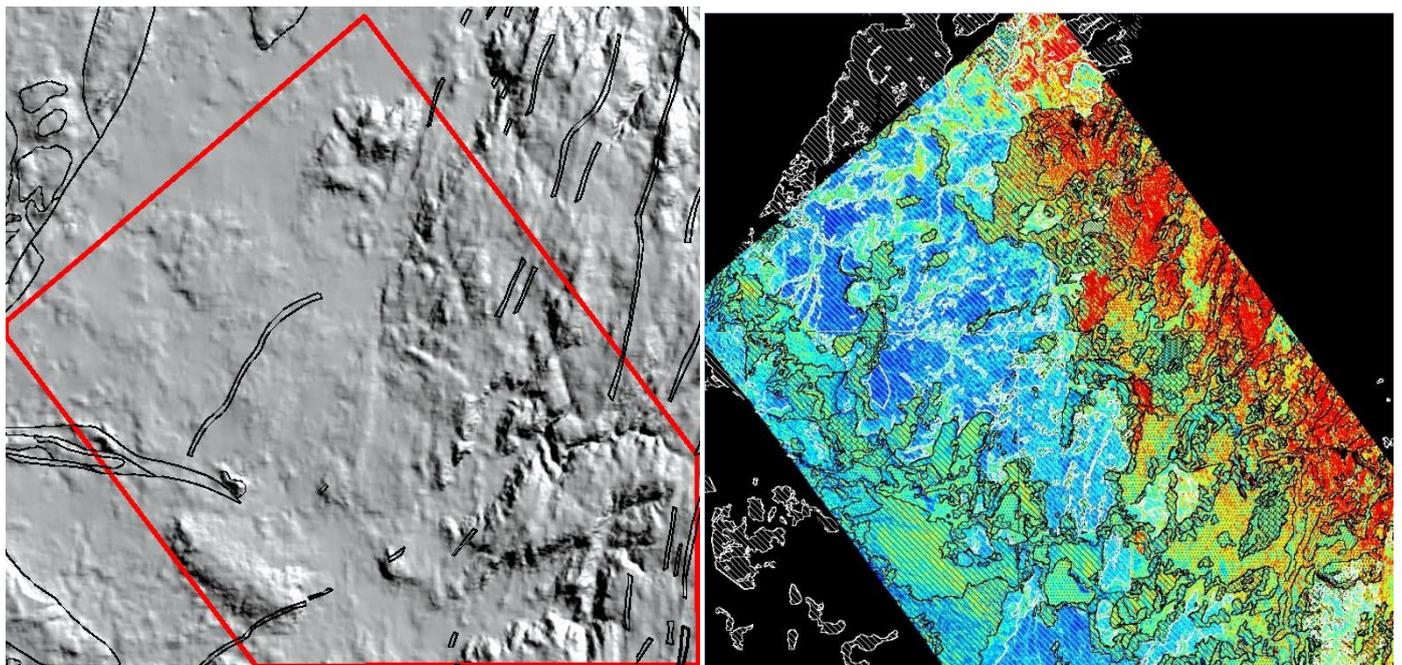
### 3.2.1 TS1340 Northern study area

The northern portion of flight line TS1340 proved an useful area for examining the response of the multi -wavelength and -polarisation radar to discriminate vegetation where limited geological outcrop and diversity of units appeared apparent by the published 1:250000 geology (Figure 14a). A range of temperate rainforest and vegetation occupies this area from the steep hills and gullies to swampy floodplains (Figure 14b). Figure 14c highlights this topography and limited geological units mapped within this lowland dominated terrain. The predominant mapped unit is the Cowrie Siltstone intruded by with mostly narrow north-easterly trending Tayatea Dyke Swarm dolerite bodies. There is a significant variation in vegetation from swampy moorland grasses to wet rainforest to open woodland (Figure 14d). Figure 14d also indicates the short wavelength and vertically polarised Cvv band exhibits an exaggerated bias in backscatter related to lower incident angles (*e.g.* compare with Figure 10).



a)

b)



c)

d)

Figure 14 a) Published 1:250000 geology of the vegetation study area; b) 1:100000 topographic map; c) shaded SRTM 25m DEM with 1:250000 geology overlay; d) Pseudo coloured Cvv backscatter response (blue:low backscatter return; red high backscatter) with 1:25,000 vegetation overlay. Black hatch: Eucalyptus obliqua wet forest (WOU), white hatch : Buttongrass moorland (MBU); black stipple : Nothofagus-Phyllocladus short rainforest (rms); black cross hatch : Eucalyptus nitida dry forest and woodland (DNI).

The Lhv band and polarisation (Figure 15) and Phv, showed a better comparison between the 1:25000 vegetation classification than the Cvv band (Figure 14d). There appeared less lower incident angle bias with this longer wavelength cross polarised bands, producing similar backscatter responses for the same vegetation class throughout the imagery. For example the Buttongrass moorland is indicated by a consistently low backscatter return (Figure 15).

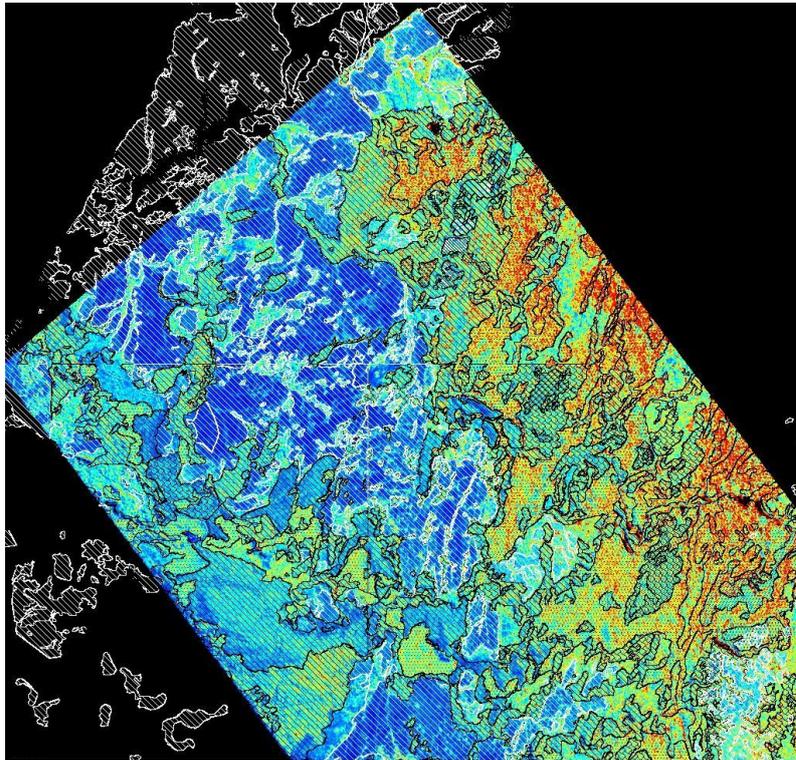


Figure 15 Pseudo coloured Lhv backscatter response (blue:low backscatter return; red high backscatter) with 1:25000 vegetation overlay. Black hatch: Eucalyptus obliqua wet forest (WOU), white hatch : Buttongrass moorland (MBU); black stipple : Nothofagus-Phyllocladus short rainforest (rms); black cross hatch : Eucalyptus nitida dry forest and woodland (DNI).

A closer comparison of the higher resolution TOPSAR DEM with SRTM 25m DEM and the published geology reveals that subtle extensions to the north easterly trending Tayatea Dyke Swarm dolerite bodies (Figure 16). The improvement in spatial resolution from the 25 metre SRTM (Figure 14c) to the 5 metre airborne TOPSAR (Figure 16) appears to significantly enhance the potential geological interpretation.

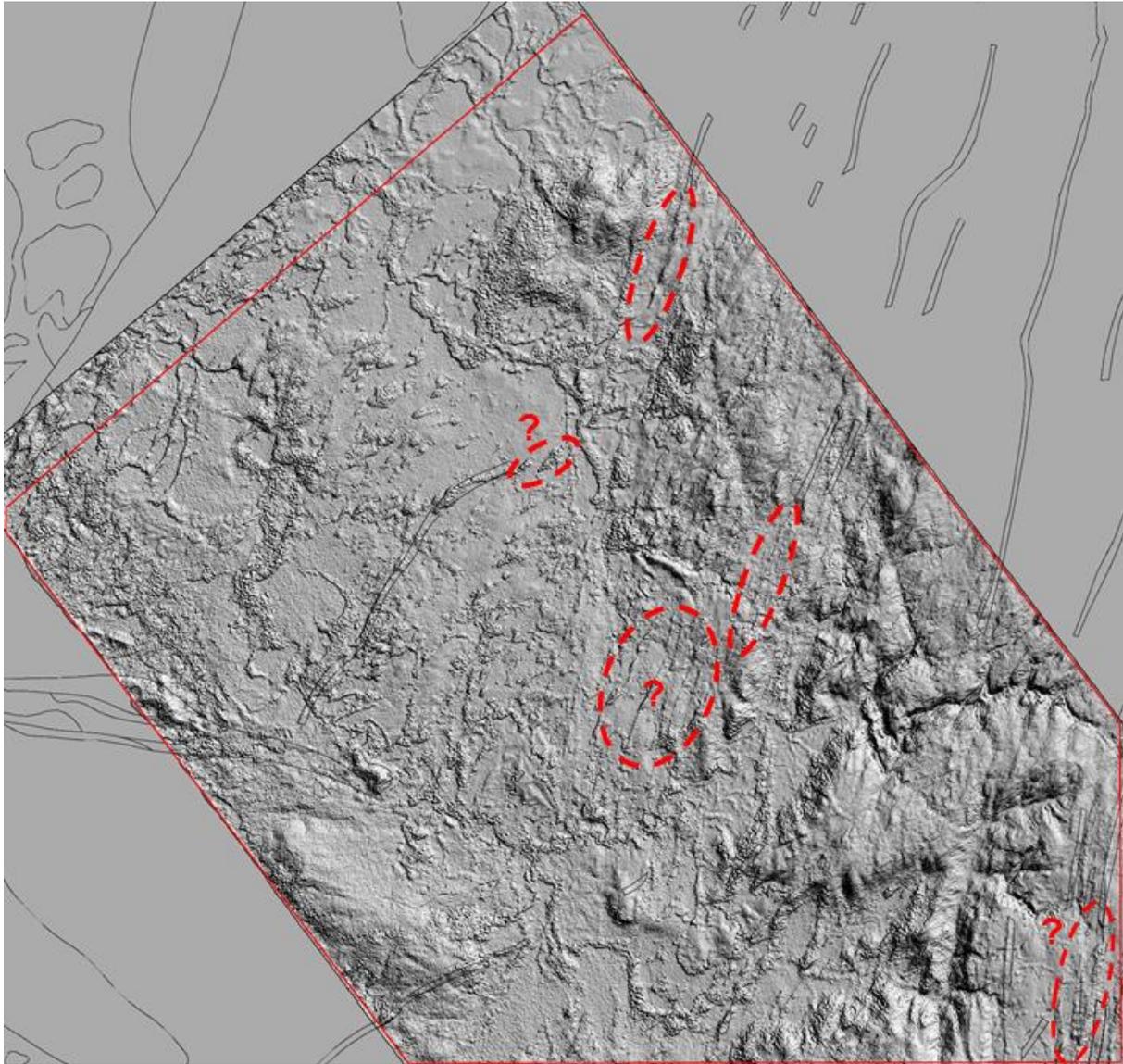


Figure 16 Artificially shaded TOPSAR DEM (45 deg. NE azimuth, 60 deg. elevation) with 1:250000 published geology. Suggested extensions or presence of dolerite dykes are highlighted within dashed red ellipses.

### 3.2.2 TS1340 Southern study area

The 1:25,000 topographic mapping of the southern TS1340 area shows a more deeply incised relief with several watercourses including the Donaldson River (Figure 17a). The published 1:25,000 geology indicates a mixture of Cowrie Siltstone, metasedimentary units, and dolerite dyke intrusives overlain by Tertiary basalts (yellow) (Figure 17b). The radiometrics imagery appears only useful for highlighting the Cowrie Siltstone and Balfour Subgroup (Figure 17c) while the aeromagnetics shows the importance of the north and northeast trending structures and boundaries (Figure 17d).

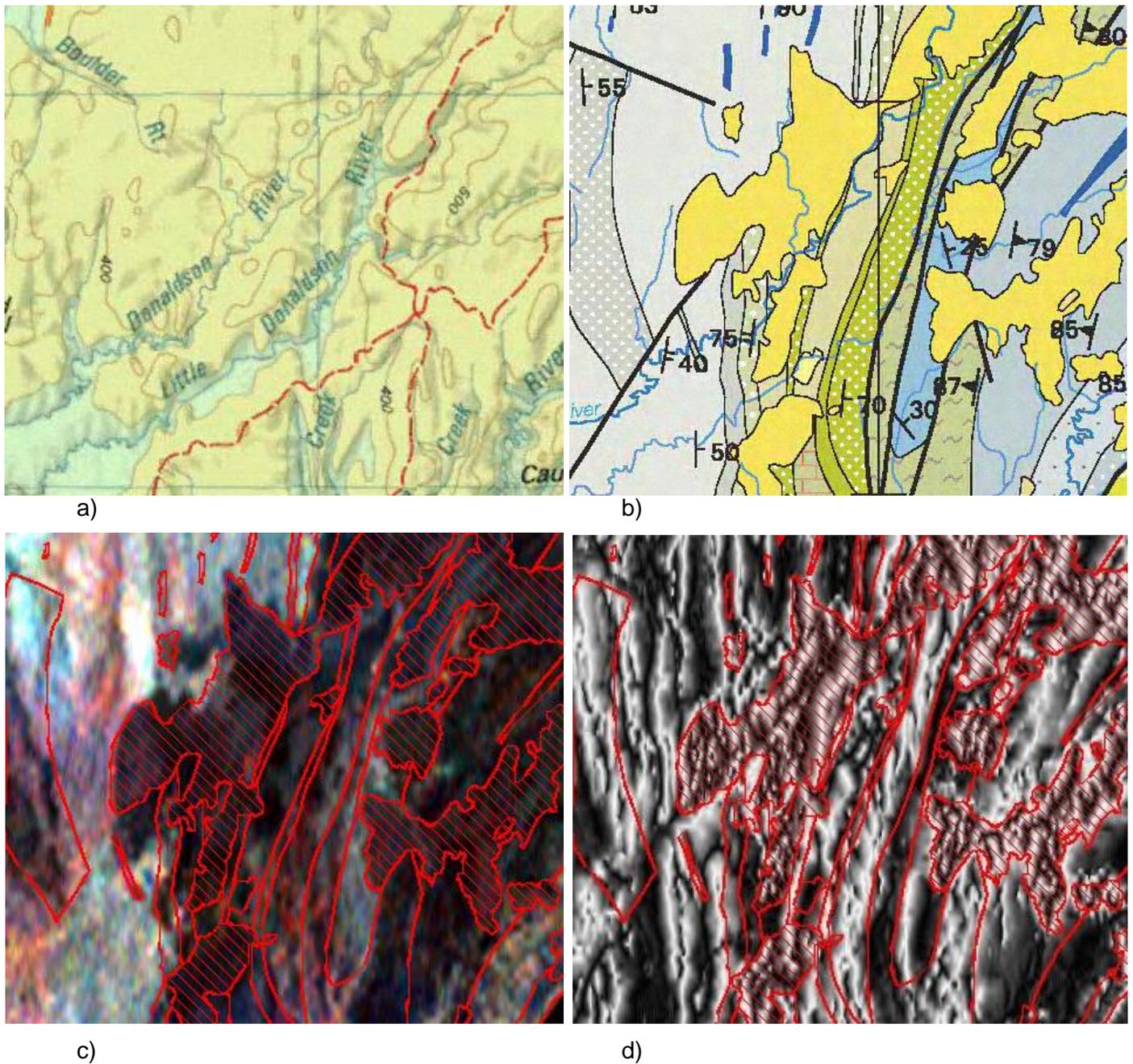


Figure 17 a) 1:25,000 topographic map of TS1340 southern study area as shown in Figure 13; b) 1:250,000 published geology with Tertiary basalt highlighted in yellow; c) Radiometric Ternary image (RGB : KThU) with geology overlay. Tertiary basalt highlighted within red hatched area; d) High frequency Tilt filtered aeromagnetics. Tertiary basalt highlighted within red hatched area.

The TOPSAR DEM product (Figure 18b) again shows a significant improvement over the SRTM 25 DEM (Figure 18a). Subtle north trending elevation features appear to be associated with dolerite dykes and their possible extension (Figure 18b).

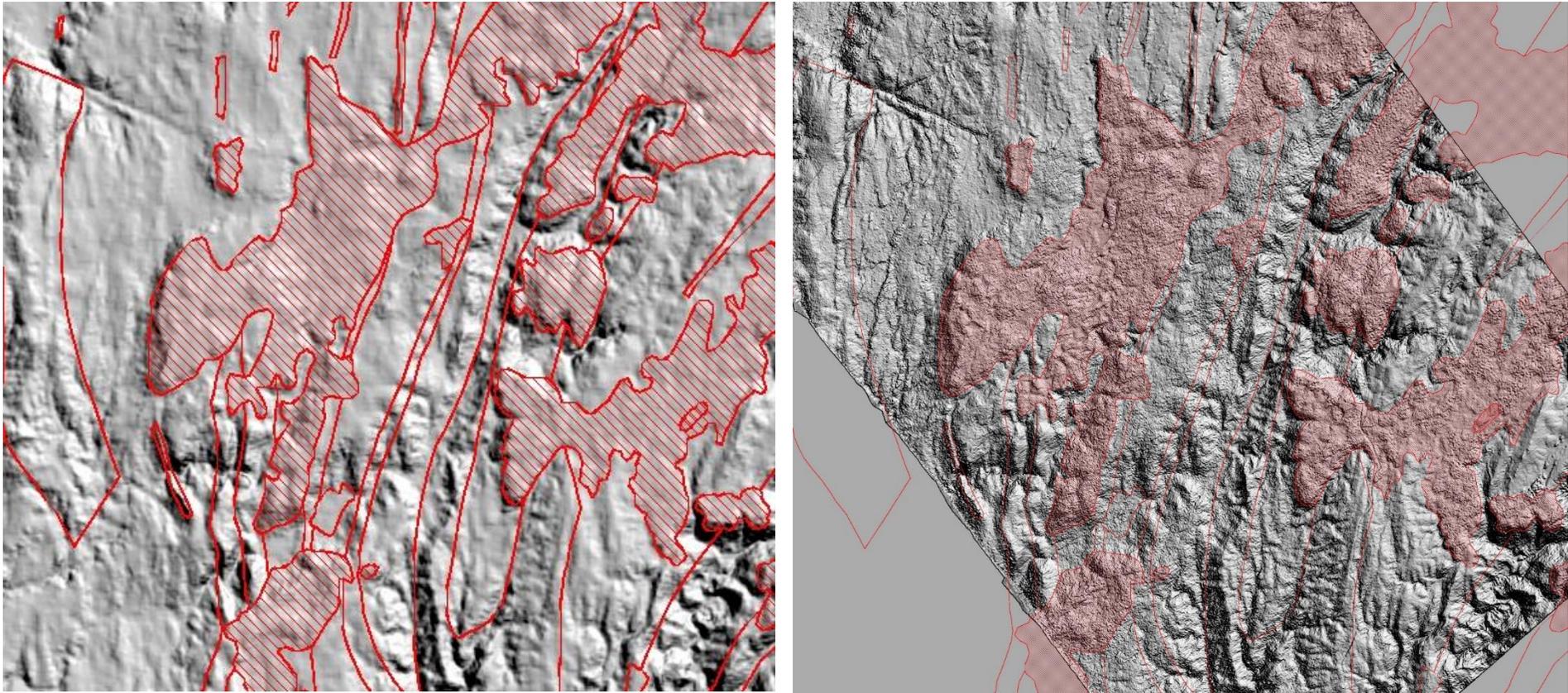


Figure 18 a) Shaded SRTM 25m DEM (45 deg. NE azimuth, 60 deg. elevation) with 1:25,000 geology overlay. Tertiary basalt highlighted within red hatched area. b) Shaded TOPSAR DEM (45 deg. NE azimuth, 60 deg. elevation) with 1:25,000 published geology overlay. Tertiary basalt highlighted within red hatched area.

An examination of the various multi-wavelength and -polarimetric bands with the published geology indicated that optimal results were obtained with Phv imagery (Figure 19). A north westerly fault line transecting the Balfour Subgroup sedimentary units (“A”) was clear within the DEMs and the Phv imagery (Figure 19). There also appeared to be a Phv anomalous area encompassing and larger than the presently mapped dolerite Tayatea Dyke nearby (“B”) although this response wasn’t replicated elsewhere (Figure 19). The overlying Tertiary basalt shows a consistently moderate to high Phv backscatter within areas shown by hatching (“C”) (Figure 19). The Arthur Metamorphic Complex unit appears to have the highest surface roughness related backscatter (“D”) while the Cowrie Siltstone shows a generally low backscatter (“E”) (Figure 19). The same area also has a diversity of vegetation types (Figure 20). Anomalous areas, “B” and “D” shown in Figure 19, also appear to have classified vegetation boundaries associated with them (Figure 20). It is possible that geobotanical associations coexist with these geological units. Further ground truthing which assist an understanding of such geological-vegetation spatial relationships.

RGB colour composite imagery combining the various bands and polarisations did not generally produce imagery that clearly discriminated either vegetation or geology. A RGB image consisting of Phh, Pvv and Phv showed some subtle variations in hue across the various units (Figure 21). The greatest benefit appeared to be the highlighting of topographic features, possibly related to geological structure or lineaments (Figure 21).

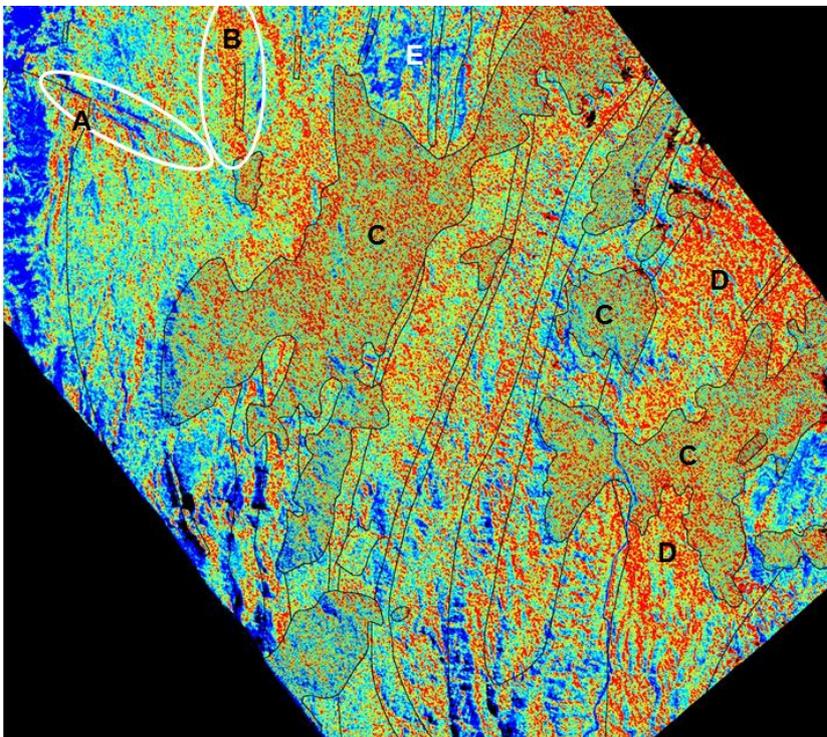


Figure 19 Pseudo coloured Phv backscatter response (blue:low backscatter return; red high backscatter) with 1:250000 geology overlay. Tertiary basalt highlighted within hatched area.

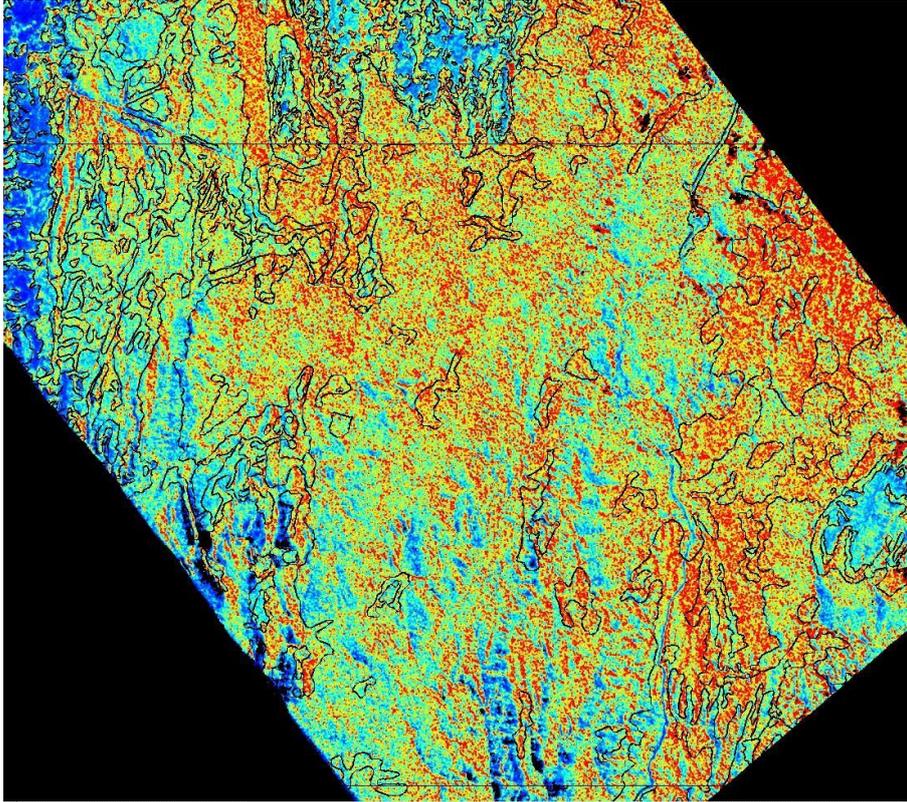


Figure 20 Pseudo coloured Phv backscatter response (blue:low backscatter return; red high backscatter) with 1:25,000 vegetation overlay.

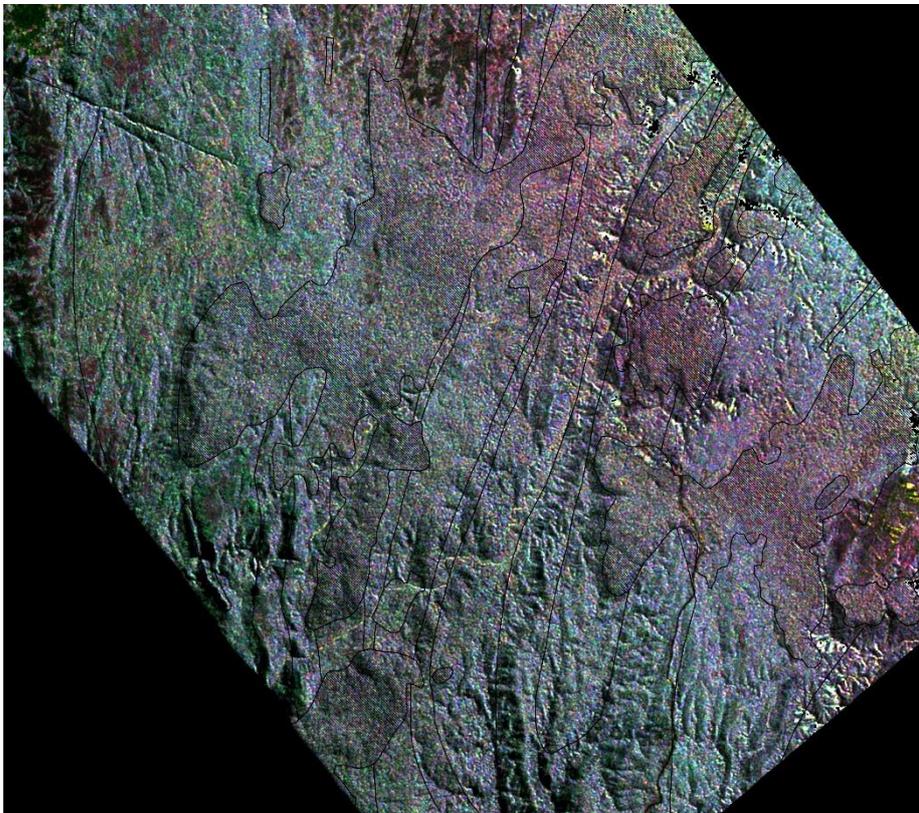
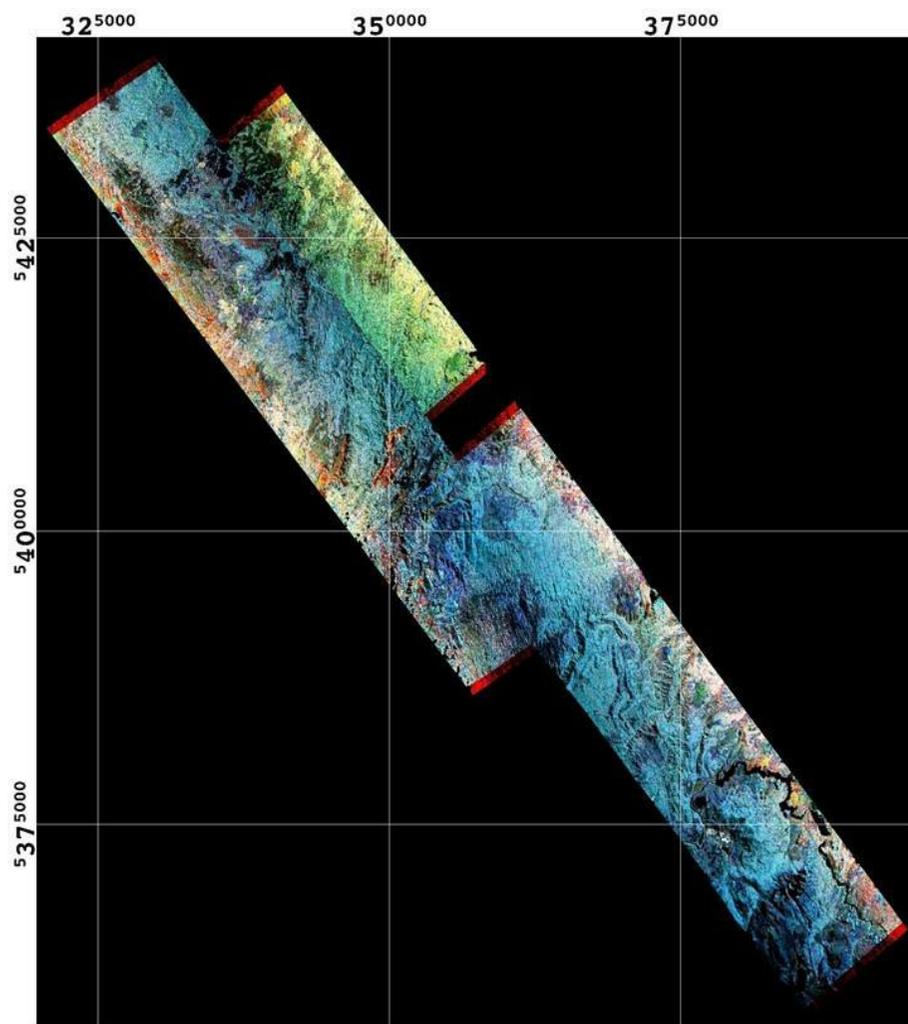


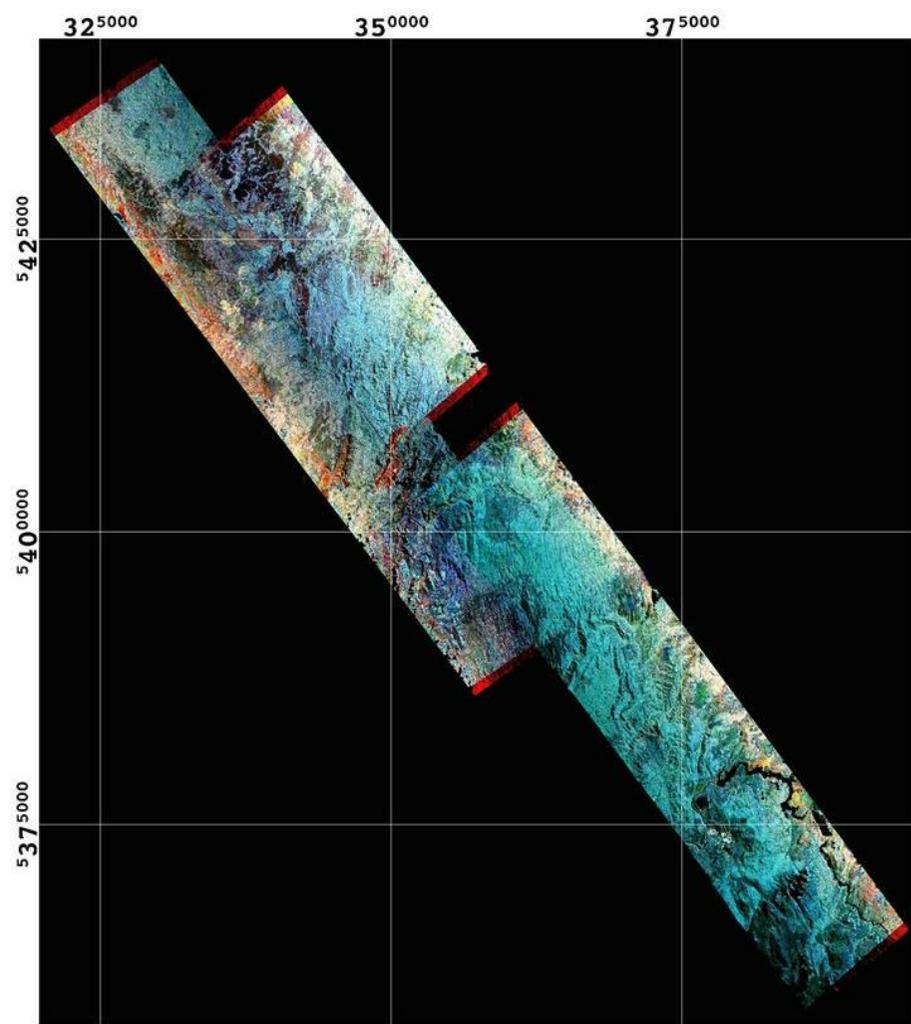
Figure 21 RGB composite image of Phh, Pvv, Phv backscatter response with 1:25,000 geology overlay.

### 3.2.3 Mosaicked polarimetric imagery

Mosaicked results combining the three flight lines of multi-wavelength and -polarimetric data were affected by the issues described in Section 2.3.4. RGB colour composite products for the combined C<sub>vv</sub>, L<sub>tp</sub> and P<sub>tp</sub> products showed that a uniform linear stretch for the three acquisitions produced such boundary issues (Figure 22a). However the application of independently calculated 99 percent linear stretching for each acquisition generally improved the mosaicked product (Figure 22b). This effect was most notable for bands such as the C<sub>vv</sub> (Figure 23a). An improvement in mosaicked results were generated using P<sub>hv</sub> and gave similar results whether an uniform or independent flight line linear stretching was applied (Figure 23b). This is consistent with the description earlier that the cross-polarized HV scattering bands are less dependent on incidence angle (and therefore, terrain), than the co-polarized (HH and VV) backscatter (Tapley, 2002).



a)



b)

Figure 22 a) RGB composite of Cvv, Ltp and Ptp with uniform linear stretch applied to each band of combined TOPSAR acquisitions; b) RGB composite of Cvv, Ltp and Ptp with 99.0% linear stretch applied independently to each band of TS1297, TS1339 and TS1340. All imagery is registered to GDA94 MGA Zone 55.

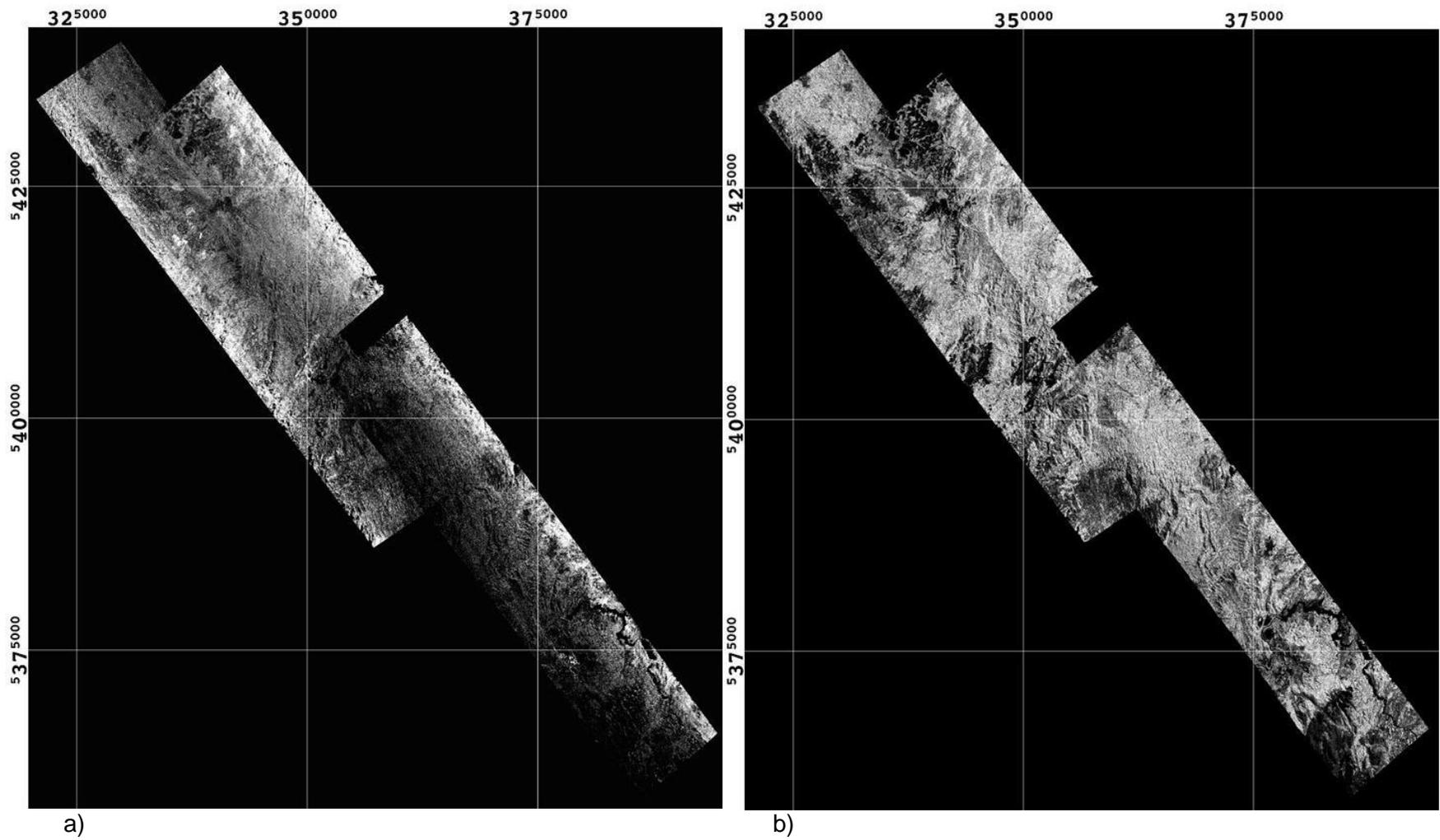


Figure 23 a) Greyscale imagery of Cvv with 99.0% linear stretch applied independently to each band of TS1297, TS1339 and TS1340; b) Greyscale imagery of Phv with uniform linear stretch applied to combined TOPSAR Phv. All imagery is registered to GDA94 MGA Zone 55.

## 4 Conclusions

This study demonstrated the successful processing and application of TOPSAR DEM and associated multi –wavelength and –polarimetric data for geoscience information in western Tasmania. In particular, a preliminary examination of the higher resolution DEM product showed its potential as a superior tool for topographic mapping and geological interpretation than available SRTM 25 imagery. A comparison of the TOPSAR DEM product with the 1:25,000 Tasmanian topographic mapping and other data sets revealed a high level of spatial and elevation accuracy. The DEM accuracy of the airborne TOPSAR data also appeared to provide consistent results, independent of the underlying geology. Although the terrain is generally forested with a variety of vegetation types, P band with a HV polarisation shows potential to map a range of surface roughness that may be related to geological units and their associated regolith. However it is recommended that the detailed 1:25,000 vegetation classification be used with any interpretation with this radar data. The complexity of radar scattering within rainforest canopies, vegetated understories and a damp regolith is likely to generate, at least partially, non-unique solutions to the question of geological versus vegetation controls. Also, an understanding of potential geobotanical associations and geologically defined variations in surface roughness would be further assisted with field work and ground truthing. Further processing of the TOPSAR data for estimates of effective layover distortion and radar shadowing using DEM information would also assist the interpretation.

## 5 Acknowledgements

This study has been kindly supported and funded by Mineral Resources Tasmania (MRT), with a considerable help from MRT's David Green. Professor Simon Jones of the School of Mathematical & Geospatial Sciences within RMIT University has also been supportive of this project. Technical assistance by Phil Wilkes, PhD candidate, has been forthcoming particularly in regard to handling LiDar data sets. Software support from Laurie Buxton of RMIT University was also helpful during this study.

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## 7 Appendix 1

### Registration ground control points of TOPSAR imagery

Note : all imagery was registered to GDA94 MGA Zone 55.

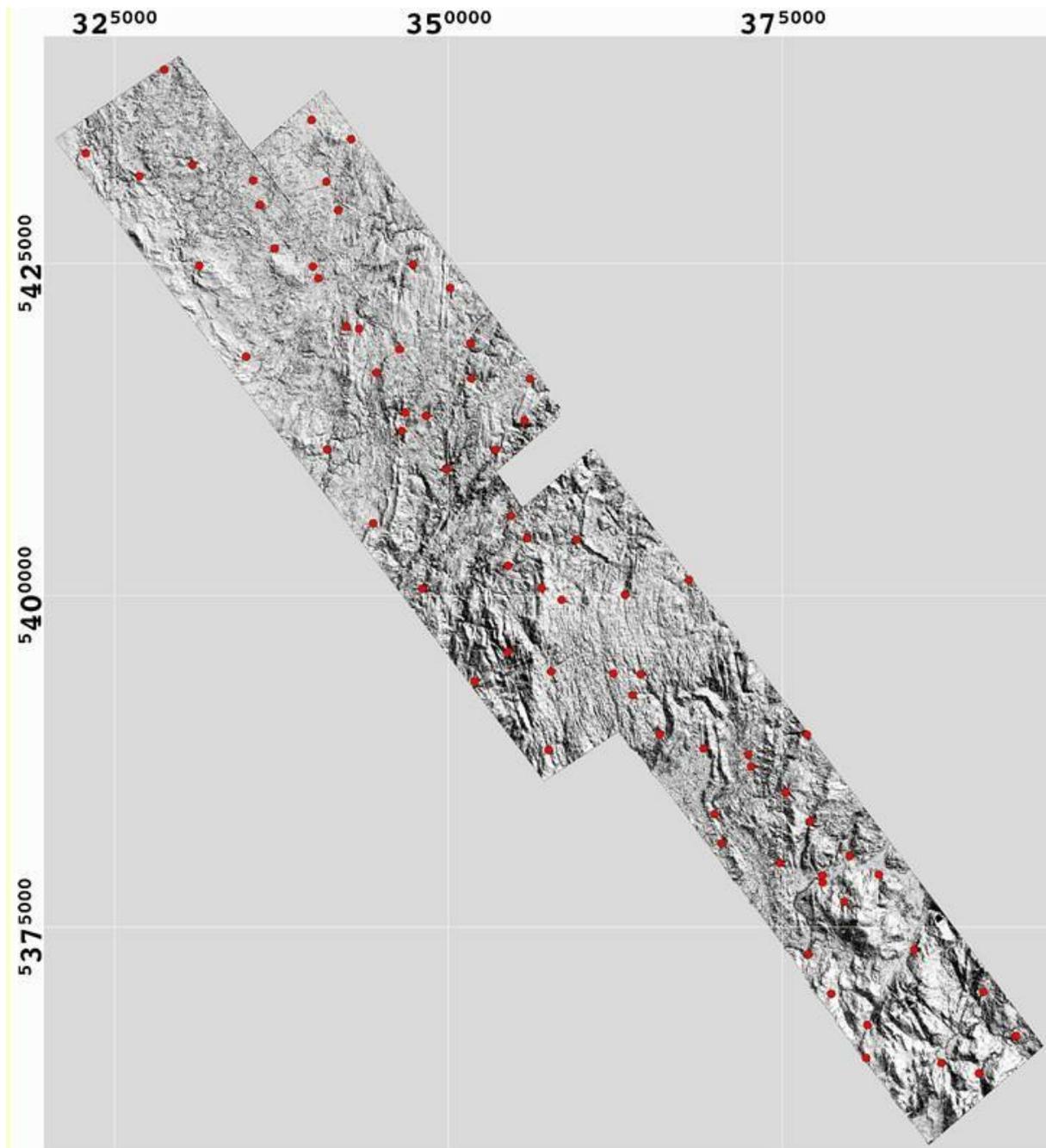


Figure A1 Shaded TOPSAR DEM with location of ground control points used for registration of each of the three TOPSAR flight lines, TS 1297, 1339 and 1340.

**TS1297 :**

```
; ENVI Image to Map GCP File  
; projection info = {Geographic Lat/Lon, WGS-84, units=Degrees}  
; warp file: C:\Tassie_MRT_AIRSAR-TOPSAR\ts1297\ts1297_c_dem_bad-value-remove  
; Map (x,y), Image (x,y)  
;
```

145.73019400	-41.85010100	0.000000	0.000000
145.61782800	-41.91090400	2284.000000	1.000000
145.21191400	-41.49717700	2284.000000	11431.000000
145.32418800	-41.43636300	0.000000	11431.000000

```
; ENVI Image to Map GCP File  
; projection info = {3, 6378137.0, 6356752.3, 0.000000, 147.000000, 500000.0, 10000000.0,  
0.999600, Geocentric Datum of Australia 1994, Map Grid of Australia (MGA 94) Zone 55,  
units=Meters}  
; warp file: G:\Tassie_MRT_AIRSAR-TOPSAR\ts1297\ts1297_c_dem_bad-value-  
remove_MGA55_RST_shade45Az60EI  
; Map (x,y), Image (x,y)  
;
```

354469.7916	5402228.9750	621.000000	1920.000000
364394.7916	5394028.9750	2688.000000	3540.000000
377069.7916	5382953.9750	5305.000000	5721.000000
376819.7916	5389503.9750	5263.000000	4404.000000
372644.7916	5387078.9750	4390.000000	4886.000000
370444.7916	5381303.9750	3940.000000	6062.000000
359594.7916	5404203.9750	1698.000000	1516.000000
379619.7916	5376903.9750	5838.000000	6915.000000
382194.7916	5378953.9750	6365.000000	6500.000000
384844.7916	5373278.9750	6923.000000	7625.000000
363219.7916	5400028.9750	2440.000000	2332.000000
365819.7916	5389528.9750	2980.000000	4435.000000
369094.7916	5388453.9750	3655.000000	4638.000000
375244.7916	5385153.9750	4931.000000	5288.000000
380019.7916	5380378.9750	5911.000000	6232.000000
374794.7916	5379803.9750	4839.000000	6351.000000
377994.7916	5378353.9750	5502.000000	6630.000000
378644.7916	5369953.9750	5648.000000	8301.000000
381344.7916	5367628.9750	6189.000000	8776.000000
390019.7916	5370103.9750	7990.000000	8263.000000
392469.7916	5366803.9750	8480.000000	8907.000000
389694.7916	5363978.9750	7915.000000	9478.000000
381244.7916	5365153.9750	6175.000000	9257.000000
386894.7916	5364753.9750	7329.000000	9310.000000
356969.7916	5400528.9750	1139.000000	2255.000000
362344.7916	5394103.9750	2251.000000	3521.000000
367994.7916	5401178.9750	3430.000000	2089.000000
377977.3770	5378879.3262	5494.000000	6524.000000
376917.0650	5372954.3391	5271.000000	7703.000000
355903.0313	5404300.5233	903.000000	1494.000000
372451.9260	5388001.5225	4345.000000	4708.000000
369902.5303	5383505.9756	3820.000000	5614.000000

Note : Root Mean Square (rms) error of registration : 9.4 pixels

**TS1339 :**

```
; ENVI Image to Map GCP File
; projection info = {3, 6378137.0, 6356752.3, 0.000000, 147.000000, 500000.0, 10000000.0,
0.999600, Geocentric Datum of Australia 1994, Map Grid of Australia (MGA 94) Zone 55,
units=Meters}
; warp file: C:\Tassie_MRT_AIRSAR-
TOPSAR\ts1339\ts1339_c_demi2_convert_DEM_masked
; Map (x,y), Image (x,y)
;
320409.6464    5434507.1693    0.000000    0.000000
329478.6014    5441404.1975    2288.000000    1.000000
366965.1318    5392380.5798    2288.000000    12273.000000
357936.3988    5385527.6382    1.000000    12273.000000
```

```
; ENVI Image to Map GCP File
; projection info = {3, 6378137.0, 6356752.3, 0.000000, 147.000000, 500000.0, 10000000.0,
0.999600, Geocentric Datum of Australia 1994, Map Grid of Australia (MGA 94) Zone 55,
units=Meters}
; warp file: G:\Tassie_MRT_AIRSAR-
TOPSAR\ts1339\ts1339_c_demi2_convert_DEM_masked_MGA55_RST_Shade135Az60EI
; Map (x,y), Image (x,y)
;
322869.7916    5433303.9750    511.000000    1748.000000
330869.7916    5432428.9750    2150.000000    1927.000000
328744.7916    5439603.9750    1703.000000    504.000000
331369.7916    5424828.9750    2276.120000    3448.160000
334870.0000    5417978.9750    2994.000000    4808.000000
342344.7916    5420228.9750    4526.000000    4350.000000
344616.4647    5416774.9417    4987.000000    5044.000000
340919.7916    5410978.9750    4255.000000    6207.000000
344394.7916    5405428.9750    4978.000000    7321.000000
349869.7916    5409503.9750    6078.000000    6498.000000
348069.7916    5400478.9750    5731.000000    8319.000000
354694.7916    5405978.9750    7073.000000    7223.000000
351969.7916    5393503.9750    6538.000000    9723.000000
358469.7916    5399653.9750    7863.000000    8486.000000
354469.7916    5395753.9750    7039.000000    9268.000000
357494.7916    5388328.9750    7670.000000    10752.000000
363794.7916    5392478.9750    8939.000000    9925.000000
357694.7916    5394253.9750    7703.000000    9563.000000
346519.7916    5412378.9750    5395.360000    5933.160000
335925.4663    5429417.0885    3188.640000    2522.540000
339866.1821    5424782.7668    4011.000000    3447.000000
326893.2004    5431585.2948    1337.000000    2089.000000
```

Note : Root Mean Square (rms) error of registration : 7.3 pixels

**TS1340 :**

```
; ENVI Image to Map GCP File
; projection info = {Geographic Lat/Lon, WGS-84, units=Degrees}
; warp file: C:\Tassie_MRT_AIRSAR-TOPSAR\ts1340\ts1340_c-demi2_convert
; Map (x,y), Image (x,y)
;
145.31069900  -41.41519900    1.000000    1.000000
145.20140100  -41.47602800    2286.000000    1.000000
144.98007200  -41.25043100    2286.000000    6234.000000
145.08926400  -41.18959000    1.000000    6234.000000
```

```
; ENVI Image to Map GCP File
; projection info = {3, 6378137.0, 6356752.3, 0.000000, 147.000000, 500000.0, 10000000.0,
0.999600, Geocentric Datum of Australia 1994, Map Grid of Australia (MGA 94) Zone 55,
units=Meters}
; warp file: C:\Tassie_MRT_AIRSAR-TOPSAR\ts1340\ts1340_c-demi2_convert_DEM-
masked_MGA55_Shade_60EI135Az
; Map (x,y), Image (x,y)
;
343319.7916  5420078.9750    2505.000000    3678.000000
347344.7916  5424928.9750    3300.000000    2704.000000
351644.7916  5418978.9750    4203.000000    3873.000000
349844.7916  5409553.9750    3875.000000    5757.000000
353519.7916  5410953.9750    4612.000000    5479.000000
355694.7916  5413178.9750    5045.000000    5033.000000
335929.0870  5429415.7599     964.000000    1830.000000
342740.8529  5434363.8219    2329.000000     824.000000
339771.1342  5435805.9494    1724.000000     545.000000
340282.2807  5423896.7677    1871.000000    2919.000000
341771.3644  5429013.9942    2155.000000    1893.000000
350144.7916  5423153.9750    3877.000000    3055.000000
348336.0057  5413514.2958    3547.000000    4971.000000
351723.9351  5416330.8849    4221.000000    4401.000000
356094.7916  5416303.9750    5114.000000    4402.000000
346769.7916  5413753.9750    3234.000000    4936.000000
335413.7162  5431249.0803     846.000000    1456.000000
336999.9600  5426133.9684    1190.170000    2478.550000
346344.7916  5418528.9750    3121.000000    3980.000000
340872.4542  5431167.4678    1962.000000    1466.000000
```

Note : Root Mean Square (rms) error of registration : 4.8 pixels

## 7 Appendix 2

### Statistical comparison between DEM results from TOPSAR acquisitions

TS1339 vs TS1340 :

STATISTICS FOR DATASET: Tas\_TOPSAR\_DEM-1339\_1340\_intersected.ers

REGION: All

	<b>TS1339</b>	<b>TS1340</b>
	-----	-----
<b>NullCells</b>	104081846	104081846
<b>Non-NullCells</b>	5164904	5164904
<b>AreaInHectares</b>	12912.26	12912.26
<b>AreaInAcres</b>	31906.892	31906.892
<b>Minimum</b>	71.364	49.9
<b>Maximum</b>	555.983	719.514
<b>Mean</b>	319.281	320.428
<b>Median</b>	311.781	311.468
<b>Std.Dev.</b>	92.952	92.556
<b>Std.Dev.(n-1)</b>	92.952	92.556
<b>Corr.Eigenval.</b>	1.991	0.009
<b>Cov.Eigenval.</b>	17131.032	75.733
<b>CorrelationMatrix</b>	TS1339	TS1340
-----	-----	-----
<b>TS1339</b>	1	0.991
<b>TS1340</b>	0.991	1
<b>Determinant</b>	0.018	
<b>Corr.Eigenvectors</b>	PC1	PC2
-----	-----	-----
<b>TS1339</b>	0.707	0.707
<b>TS1340</b>	0.707	-0.707
<b>Inv.ofCorr.Ev.</b>	PC1	PC2
-----	-----	-----
<b>TS1339</b>	0.707	0.707
<b>TS1340</b>	0.707	-0.707

<b>CovarianceMatrix</b>	TS1339	TS1340
-----	----	----
TS1339	8640.139	8527.57
TS1340	8527.57	8566.626
<b>Determinant</b>	1297392.466	

<b>Cov.EigenVectors</b>	PC1	PC2
-----	----	----
TS1339	0.709	0.706
TS1340	0.706	-0.709

<b>Inv.ofCov.Ev.</b>	PC1	PC2
-----	----	----
TS1339	0.709	0.706
TS1340	0.706	-0.709

TS1297 vs TS1339 :

STATISTICS FOR DATASET: Tas\_TOPSAR\_DEM-1297-1339\_intersected.ers  
REGION:All

	<b>TS1297</b>	<b>TS1339</b>
	----	----
<b>NullCells</b>	250922347	250922347
<b>Non-NullCells</b>	3200488	3200488
<b>AreaInHectares</b>	8001.43	8001.43
<b>AreaInAcres</b>	19771.965	19771.965
<b>Minimum</b>	4.619	128.584
<b>Maximum</b>	833.887	834.257
<b>Mean</b>	505.289	501.541
<b>Median</b>	542.347	542.064
<b>Std.Dev.</b>	132.846	135.511
<b>Std.Dev.(n-1)</b>	132.846	135.511
<b>Corr.Eigenval.</b>	1.995	0.005
<b>Cov.Eigenval.</b>	35916.763	94.498

<b>CorrelationMatrix</b>	TS1297	TS1339
-----	----	----
TS1297	1	0.995
TS1339	0.995	1
<b>Determinant</b>	0.01	

<b>Corr.EigenVectors</b>	PC1	PC2
-----	----	----
TS1297	0.707	0.707

<b>TS1339</b>	0.707	-0.707
---------------	-------	--------

<b>Inv.ofCorr.Ev.</b>	PC1	PC2
-----	----	----
<b>TS1297</b>	0.707	0.707
<b>TS1339</b>	0.707	-0.707

<b>CovarianceMatrix</b>	TS1297	TS1339
-----	----	----
<b>TS1297</b>	17648.111	17907.564
<b>TS1339</b>	17907.564	18363.15
<b>Determinant</b>	3394068.61	

<b>Cov.Eigenvectors</b>	PC1	PC2
-----	----	----
<b>TS1297</b>	0.7	-0.714
<b>TS1339</b>	0.714	0.7

<b>Inv.ofCov.Ev.</b>	PC1	PC2
-----	----	----
<b>TS1297</b>	0.7	0.714
<b>TS1339</b>	-0.714	0.7

## 7 Appendix 3

### Statistical comparison between DEM results from TOPSAR (TS1297), SRTM25m, Photogrammetric 25m and LiDar (Meredith)

STATISTICS FOR DATASET: Tas\_DEM\_analysis4b.ers

REGION: Meredith Granite

	<b>TOPSAR</b>	<b>SRTM25m</b>	<b>Photo25m</b>	<b>LiDAR</b>
	----	----	----	----
<b>Null cells</b>	75923	75923	75923	75923
<b>Non-Null cells</b>	70578035	70578035	70578035	70578035
<b>Area(Ha)</b>	7057.863	7057.863	7057.863	7057.863
<b>Area(acres)</b>	17440.361	17440.361	17440.361	17440.361
<b>Minimum</b>	309.724	322.547	313.056	312.168
<b>Maximum</b>	799.251	718.522	730	1160.635
<b>Mean</b>	564.214	562.759	562.079	561.317
<b>Median</b>	567.873	565.391	567.131	564.056
<b>Std.Dev</b>	48.533	48.441	51.048	51.886
<b>Std.Dev(n-1)</b>	48.533	48.441	51.048	51.886
<b>Corr.Eigenval.</b>	3.922	0.053	0.017	0.008
<b>Cov.Eigenval.</b>	9802.053	137.109	40.746	20.148

<b>CorrelationEigenVector</b>	<b>TOPSAR</b>	<b>SRTM25m</b>	<b>Photo25m</b>	<b>LiDAR</b>
-----	----	----	----	----
<b>TOPSAR</b>	1	0.984	0.984	0.953
<b>SRTM25m</b>	0.984	1	0.992	0.965
<b>Photo25m</b>	0.984	0.992	1	0.967
<b>LiDAR</b>	0.953	0.965	0.967	1
<b>Determinant</b>	0			
<b>Corr.EigenVector</b>	PC1	PC2	PC3	PC4
-----	----	----	----	----
<b>TOPSAR</b>	0.5	-0.454	0.737	-0.022
<b>SRTM25m</b>	0.502	-0.211	-0.491	-0.68
<b>Photo25m</b>	0.503	-0.173	-0.425	0.733
<b>LiDAR</b>	0.495	0.848	0.186	-0.032
<b>Inv.</b>	PC1	PC2	PC3	PC4
-----	----	----	----	----
<b>TOPSAR</b>	0.5	0.502	0.503	0.495

<b>SRTM25m</b>	-0.454	-0.211	-0.173	0.848
<b>Photo25m</b>	0.737	-0.491	-0.425	0.186
<b>LiDAR</b>	-0.022	-0.68	0.733	-0.032

<b>Covariance</b>	<b>TOPSAR</b>	<b>SRTM25m</b>	<b>Photo25m</b>	<b>LiDAR</b>
-----	----	----	----	----
<b>TOPSAR</b>	2355.497	2313.223	2438.293	2399.021
<b>SRTM25m</b>	2313.223	2346.567	2452.582	2424.755
<b>Photo25m</b>	2438.293	2452.582	2605.882	2560.855
<b>LiDAR</b>	2399.021	2424.755	2560.855	2692.11
<b>Determinant</b>	1103315978			

<b>Cov.</b>	<b>TOPSAR</b>	<b>SRTM25m</b>	<b>Photo25m</b>	<b>LiDAR</b>
-----	----	----	----	----
<b>TOPSAR</b>	0.485	-0.446	0.752	0
<b>SRTM25m</b>	0.487	-0.228	-0.449	-0.714
<b>Photo25m</b>	0.513	-0.206	-0.453	0.7
<b>LiDAR</b>	0.514	0.841	0.167	-0.022

<b>Inv.</b>	<b>PC1</b>	<b>PC2</b>	<b>PC3</b>	<b>PC4</b>
-----	----	----	----	----
<b>TOPSAR</b>	0.485	0.487	0.513	0.514
<b>SRTM25m</b>	-0.446	-0.228	-0.206	0.841
<b>Photo25m</b>	0.752	-0.449	-0.453	0.167
<b>LiDAR</b>	0	-0.714	0.7	-0.022

STATISTICS FOR DATASET: Tas DEM analysis4b no Meredith.ers  
REGION: No Meredith Granite

	<b>TOPSAR</b>	<b>SRTM25m</b>	<b>Photo25m</b>	<b>LiDAR</b>
	----	----	----	----
<b>Non-NullCells</b>	85406797	85406797	85406797	85406797
<b>AreaInHectares</b>	8540.752	8540.752	8540.752	8540.752
<b>AreaInAcres</b>	21104.66	21104.66	21104.66	21104.66
<b>Minimum</b>	3.531	105.086	97	94.982
<b>Maximum</b>	917.031	915.842	920	928.798
<b>Mean</b>	335.072	341.235	337.936	337.196
<b>Median</b>	321.115	326.777	318.824	319.721
<b>Std.Dev.</b>	160.957	159.243	160.217	159.258
<b>Std.Dev.(n-1)</b>	160.957	159.243	160.217	159.258
<b>Corr.Eigenval.</b>	3.995	0.003	0.001	0.001
<b>Cov.Eigenval.</b>	102182.685	69.109	29.725	16.677

<b>CorrelationMatrix</b>	TOPSAR	SRTM25m	Photo25m	LiDAR
-----	----	----	----	----
TOPSAR	1	0.998	0.998	0.998
SRTM25m	0.998	1	0.999	0.999
Photo25m	0.998	0.999	1	0.999
LiDAR	0.998	0.999	0.999	1
Determinant	0			
<b>Corr.Eigenvectors</b>	PC1	PC2	PC3	PC4
-----	----	----	----	----
TOPSAR	0.5	-0.822	-0.252	-0.103
SRTM25m	0.5	0.028	0.86	0.102
Photo25m	0.5	0.334	-0.384	0.701
LiDAR	0.5	0.46	-0.223	-0.699
<b>Inv.ofCorr.Ev.</b>	PC1	PC2	PC3	PC4
-----	----	----	----	----
TOPSAR	0.5	0.5	0.5	0.5
SRTM25m	-0.822	0.028	0.334	0.46
Photo25m	-0.252	0.86	-0.384	-0.223
LiDAR	-0.103	0.102	0.701	-0.699
<b>CovarianceMatrix</b>	TOPSAR	SRTM25m	Photo25m	LiDAR
-----	----	----	----	----
TOPSAR	25907.312	25587.659	25736.169	25574.251
SRTM25m	25587.659	25358.315	25484.135	25331.676
Photo25m	25736.169	25484.135	25669.439	25498.565
LiDAR	25574.251	25331.676	25498.565	25363.13
Determinant	3500760358			
<b>Cov.Eigenvectors</b>	PC1	PC2	PC3	PC4
-----	----	----	----	----
TOPSAR	0.503	-0.822	-0.248	-0.102
SRTM25m	0.498	0.033	0.86	0.105
Photo25m	0.501	0.337	-0.388	0.697
LiDAR	0.498	0.459	-0.22	-0.702
<b>Inv.ofCov.Ev.</b>	PC1	PC2	PC3	PC4
-----	----	----	----	----
TOPSAR	0.503	0.498	0.501	0.498
SRTM25m	-0.822	0.033	0.337	0.459
Photo25m	-0.248	0.86	-0.388	-0.22
LiDAR	-0.102	0.105	0.697	-0.702

REGION:All (Inc. Meredith & non-Meredith intrusives) :

	<b>TOPSAR</b>	<b>SRTM25m</b>	<b>Photo25m</b>	<b>LiDAR</b>
	----	----	----	----
<b>NullCells</b>	1206461396	1206461396	1206461396	1206461396
<b>Non-NullCells</b>	176478100	176478100	176478100	176478100
<b>ArealnHectares</b>	17647.96	17647.96	17647.96	17647.96
<b>ArealnAcres</b>	43609.061	43609.061	43609.061	43609.061
<b>Minimum</b>	3.531	102.597	97	94.982
<b>Maximum</b>	917.031	915.842	920	1160.635
<b>Mean</b>	434.5	437.099	434.686	434.036
<b>Median</b>	481.691	483.805	479.566	477.951
<b>Std.Dev.</b>	166.346	162.863	164.762	164.268
<b>Std.Dev.(n-1)</b>	166.346	162.863	164.762	164.268
<b>Corr.Eigenval.</b>	3.993	0.004	0.002	0.001
<b>Cov.Eigenval.</b>	108145.165	106.368	47.345	27.093

<b>CorrelationMatrix</b>	<b>TOPSAR</b>	<b>SRTM25m</b>	<b>Photo25m</b>	<b>LiDAR</b>
-----	----	----	----	----
<b>TOPSAR</b>	1	0.998	0.998	0.996
<b>SRTM25m</b>	0.998	1	0.999	0.998
<b>Photo25m</b>	0.998	0.999	1	0.998
<b>LiDAR</b>	0.996	0.998	0.998	1
<b>Determinant</b>	0			

<b>Corr.Eigenvectors</b>	<b>PC1</b>	<b>PC2</b>	<b>PC3</b>	<b>PC4</b>
-----	----	----	----	----
<b>TOPSAR</b>	0.5	-0.748	0.435	0.042
<b>SRTM25m</b>	0.5	-0.023	-0.55	-0.669
<b>Photo25m</b>	0.5	0.117	-0.443	0.734
<b>LiDAR</b>	0.5	0.653	0.558	-0.108

<b>Inv.ofCorr.Ev.</b>	<b>PC1</b>	<b>PC2</b>	<b>PC3</b>	<b>PC4</b>
-----	----	----	----	----
<b>TOPSAR</b>	0.5	0.5	0.5	0.5
<b>SRTM25m</b>	-0.748	-0.023	0.117	0.653
<b>Photo25m</b>	0.435	-0.55	-0.443	0.558
<b>LiDAR</b>	0.042	-0.669	0.734	-0.108

<b>CovarianceMatrix</b>	<b>TOPSAR</b>	<b>SRTM25m</b>	<b>Photo25m</b>	<b>LiDAR</b>
-----	----	----	----	----
<b>TOPSAR</b>	27670.998	27033.976	27342.512	27220.512
<b>SRTM25m</b>	27033.976	26524.266	26805.697	26696.373

<b>Photo25m</b>	27342.512	26805.697	27146.611	27016.629
<b>LiDAR</b>	27220.512	26696.373	27016.629	26984.095
<b>Determinant</b>	14755001395			

<b>Cov.Eigenvectors</b>	PC1	PC2	PC3	PC4
-----	----	----	----	----
<b>TOPSAR</b>	0.505	-0.75	0.424	0.046
<b>SRTM25m</b>	0.495	-0.014	-0.542	-0.679
<b>Photo25m</b>	0.501	0.124	-0.455	0.726
<b>LiDAR</b>	0.499	0.649	0.565	-0.101

<b>Inv.ofCov.Ev.</b>	PC1	PC2	PC3	PC4
-----	----	----	----	----
<b>TOPSAR</b>	0.505	0.495	0.501	0.499
<b>SRTM25m</b>	-0.75	-0.014	0.124	0.649
<b>Photo25m</b>	0.424	-0.542	-0.455	0.565
<b>LiDAR</b>	0.046	-0.679	0.726	-0.101